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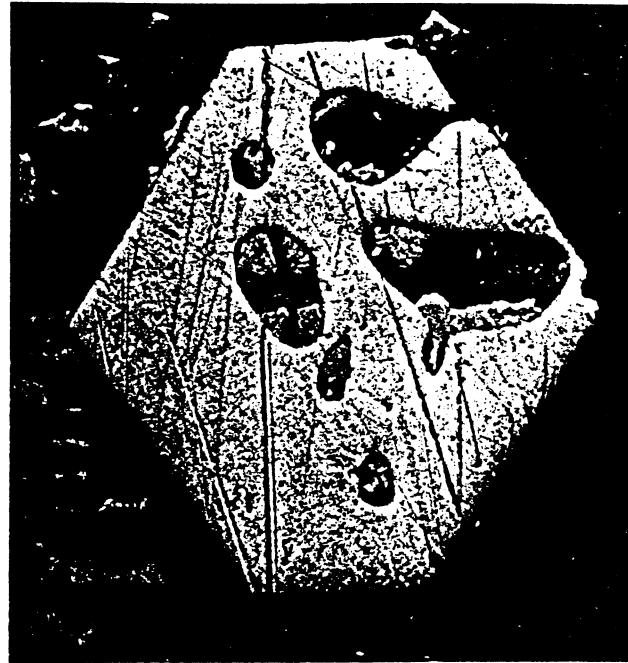
NORSK HYDRO A/S

GOLD EXPLORATION IN THE BLEKA FOLD AREA, TELEMARK.

BY

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COPENHAGEN 1984

SUMMARY.

The gold-bismuth bearing Bleka-type veins of the Bleka Fold area (170 km^2) of central Telemark have been investigated in detail. The investigation was concentrated partly in the Bleka Mine area and in the Espelid area 2-3 km north of Bleka where the hitherto largest vein swarm of the area was found and mapped at scale 1:1000. Apart from this, reconnaissance investigations of other known vein occurrences include Ny-staul, Barstad, Blengsdalen and Gjuv.

The results of this investigation strongly indicate that all of the known occurrences including the newly discovered Espelid vein swarm have a common ore genesis and a conceptual model for the veins is proposed. This concludes that the 40° - 90° striking gold-bismuth bearing quartz-tourmaline-carbonate veins were emplaced in their amphibolite host rock along ENE-WSW dextral shear faults resulting from ESE-WNW max compression during N-S to NNE-SSW folding in the Sveconorwegian orogenic period. Major ore lenses seem to be located in areas where max tension and dilation occurs e.g. merging en echelon vein segments. Pronounced hydrothermal alteration envelopes characterize the mineralization and also, together with the mineral assembly of the veins, indicate the high temperature nature of the veins. In summary, the location of the Bleka-type veins is a result of regional tectonics combined with favorable host rock lithology whereas the high temperature nature and to some degree the mineral assembly is controlled by the intrusion of underlying granite. Due to the high-grade nature of the known ore lenses (25 ppm Au and 0.5 % Bi) and due to the significant size of the Espelid vein swarm as compared with that of Bleka the economic potential is considered good here. Furthermore, a bismuth (locally gold) halo is observed in the northern end of the vein swarm probably indicative of deeper-lying gold-bismuth mineralization. In conclusion the Espelid vein swarm is considered a prime drilling target for a small to medium sized gold-bismuth deposit.

RECOMMENDATIONS.

Prime importance - the Espelid vein swarm.

- 1) It is recommended to carry out drilling on the Espelid vein swarm. It is suggested that gold-bismuth mineralization is located 200-300 m below the present surface and this should be tested in a first step by drilling of two or three holes totalling in the order of 500-1000 m from one drill location provisionally situated in the northern part of the vein swarm just south of Nystaul (see enclosed Map 3). The exact position of this locality and the details of this drilling programme would benefit from the following additional investigations:
 - a) analysis of all remaining chip samples from the area for bismuth in order to get a more precise outline of Bi-halos.
 - b) fluid-inclusion studies on quartz-(tourmaline)-(carbonate) which could indicate whether the Espelid veins are located at a 200-300 m higher level than the known gold-bearing veins of the Bleka Fold area as suggested by present day position.
 - c) a detailed microscopic investigation of the gold and bismuth minerals and their paragenetic position in the different occurrences might also help in the interpretation of relative levels.
 - d) local trenching partly along and partly across the veins in the more promising areas to identify areas of max dilation e.g. merging between en echelon vein segments.

Secondary importance - Bleka-Blengsdalen-Gjuv.

- 2) Further detailed geological investigations around Blengsdalen and Gjuv are recommended in order to be able to make a better estimate of the economic potential of these.
- 3) On a long-term basis it is suggested to investigate further the merge zone at the bleka Mine possibly by drilling. In this context the area NE of Sverveli should also be investigated.
- 4) The SW and NE continuation of the recognized gold-anomalous belt (outlined from stream-sediment survey) stretching from Espelid-Bleka-Blengsdalen to NE of Gjuv should be tested by supplementary stream-sediment sampling.

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ENCLOSED:

- MAP 1 - Structural interpretation of the Bleka Fold area -
Telemark. Based on air photo mosaic.
MAP 2 - The Bleka Vein - Telemark.
MAP 3 - The Espelid Vein Swarm.

SEPARATE EXHIBIT MAP:

- SAMPLE MAP 1 - Svartdal.
SAMPLE MAP 2 - Ambjørndalen.
SAMPLE MAP 3 - Nutheim.
SAMPLE MAP 4 - Ståldalen.
SAMPLE MAP 5 - Barstadhovi.
SAMPLE MAP 6 - Hovdejord.
SAMPLE MAP 7 - Svinom.
SAMPLE MAP 8 - Bisminuten.
SAMPLE MAP 9 - Hjartdal.
SAMPLE MAP 10 - Gvammen.
SAMPLE MAP 11 - Kivikdalen.
SAMPLE MAP 12 - "Flatdal".

1. INTRODUCTION.

The investigated area covers approximately 170 km² of the Bleka Fold in central Telemark, Norway (Fig.1).

The main aim of the project was to carry out detailed investigations on gold-bismuth mineralized veins in order to propose a mineralization model for the area.

The work was carried out during a seven-week period from late May until early July. It comprises a photogeological interpretation (lineament analysis + limited photogeological mapping) of the entire area as well as five weeks of detailed field work in selected areas.

The field work comprises:

- 1) Detailed mapping and sampling of the Bleka Mine area (0.5 km²). This includes both surface and subsurface investigations.
- 2) Detailed mapping and sampling of the Espelid area.
- 3) Investigations of other known vein occurrences.
- 4) Reconnaissance for gold-bearing veins in unknown areas.

Due to the finding of a previously unrecognized vein swarm in the Espelid area approximately half of the time was spent on detailed investigations in this area.

The field work was coordinated with stream-sediment sampling for gold in central Telemark conducted by B. Sieborg.

A total of 203 litho-samples (Appendix I) comprising 145 chip samples (1-10 kg according to vein thickness) and 58 grab samples (0.2-2 kg) were collected. Of these 193 have been analyzed for gold and 57 for one or more associated elements (Bi-Ag-Cu-Pb) (Appendix II).

Finally, two polished sections (one rock chip and one pan concentrate of ore) have been investigated ore microscopically.

We acknowledge the fruitfull discussions in the field with P.A.Lindberg, F.D.Pedersen, L.Rasmussen, B.Sieborg and T.Vrålstad.

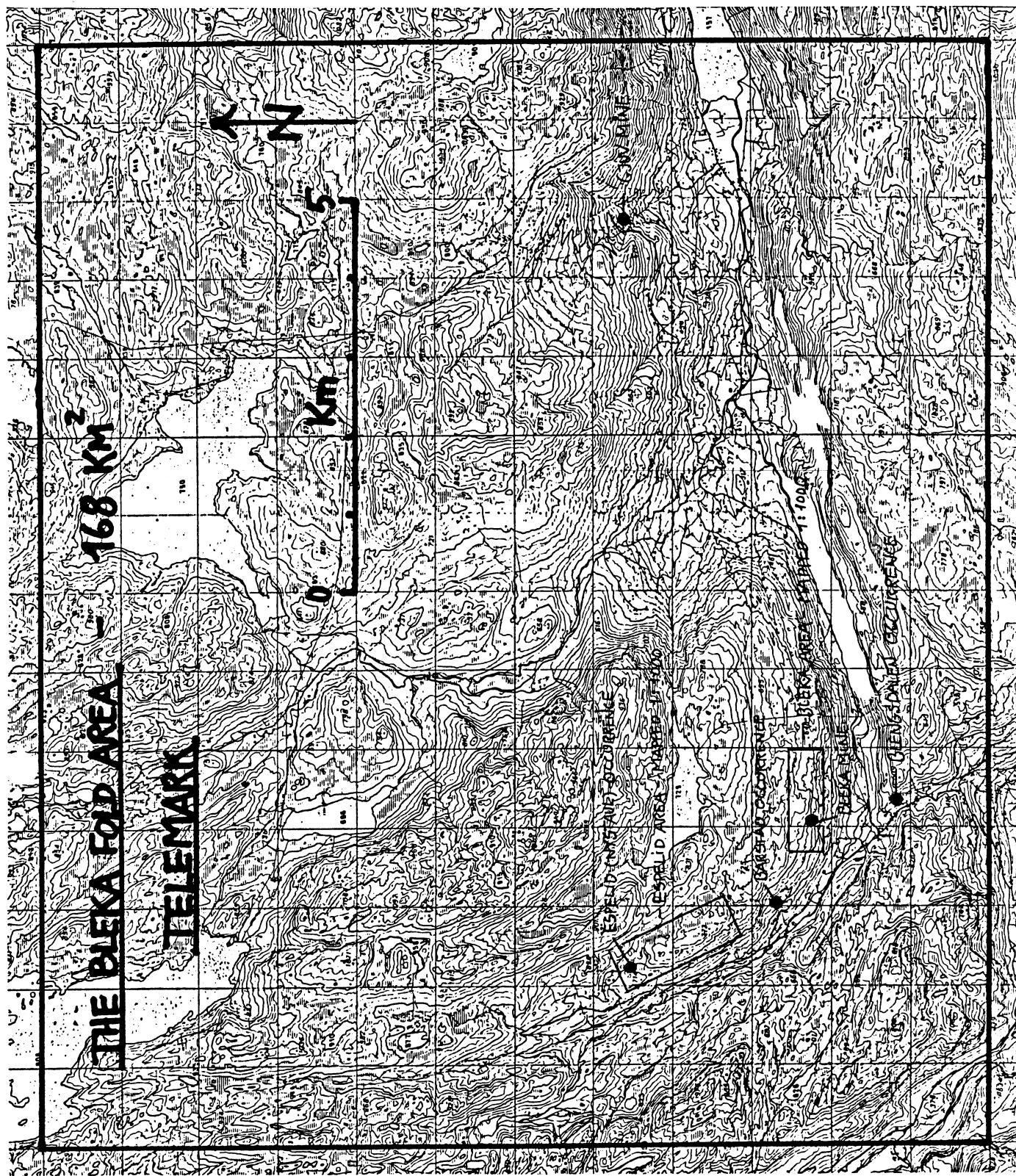


Fig.1. Location map.

2. GEOLOGICAL SETTING.

In a regional context the Bleka Fold area of central Telemark belongs to the Southern Precambrian Province of Norway. This province is bounded by the coast of Southern Norway to the south, by the Cambro-Silurian of the Oslo Region to the east and by overlying Cambrian autochton or Caledonian thrust nappes to the north and northwest. Traditionally, the province has been subdivided into : the Egersund Anorthosite Province of Rogaland in the extreme southwest, the southern basement gneisses with post-tectonic granites, the Kongsberg-Bamble Complex along the eastern margin and the Telemark Supracrustal Suite of the central area.

In a semi-regional context the geology of central Telemark is dominated by the Proterozoic supracrustals of the Telemark Supergroup which comprise an up to 7 km thick sequence of felsic and mafic volcanics, quartzites, conglomerates, schists and subordinate marbles. The supracrustals of the Telemark Supergroup have been divided into four groups : the Rjukan Group, the Seljord Group, the Bandak Group and the Heddevatn Group. The Rjukan Group (predominantly felsic and mafic volcanics - up to 1000 m thick) and the Seljord Group (predominantly quartzites, conglomerates and schists - up to 4000 m thick) have been folded and eroded before the deposition of the overlying Bandak Group (dominated by a mixed sequence of felsic and mafic volcanics and clastic sediments with subordinate calcareous sediments - up to 2200 m). The western continuation of the Bandak Group called the Blåbergås Group is conformably overlain by the Heddevatn Group (predominantly andesitic flows - up to 3200 m) which comprises the youngest deposits of the area.

Fig.2 shows a schematic representation of the stratigraphy of the Telemark Supergroup.

The last registered major geological event was the intrusion of a suite of postkinematic granites during the period 1100-850 ma. In the central Telemark area the granites are of calc-alkaline affinity.

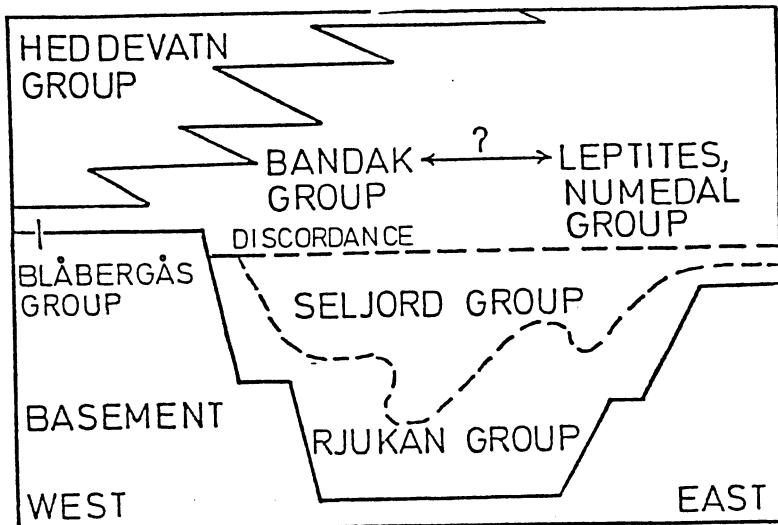


Fig.2. Schematic representation of the stratigraphy of the Telemark Supergroup. From Pedersen (1984).

A tentative scheme of the tectonic development of the central Telemark area has been proposed by Pedersen (1984):

- ~ 800 ma. 8) Drag folding by late faults
| 7) Doming around granite diapirs
| 6) Folding around N-S axes
| 5) Granitization
non-linear 4) Deposition of the Bandak and Heddevatn
time scale Groups
| 3) Uplift and erosion
| 2) Folding around E-W to ENE-WSW axes
| 1) Deposition of the Rjukan and Seljord
| Groups
~ 1650 ma.

The summary description of the regional to semi-regional geological setting as outlined above is mainly based on a recent comprehensive study by Pedersen (1984) who has interpreted the geology and proposed some conceptual models for the Southern Precambrian Province.

In a local context the geology of the investigated area of the Bleka Fold is dominated by a prominent NNE-SSW trending anti-clinal fold structure which in the core towards NNE exposes felsic volcanics of the Tuddal Formation (Rjukan Group) and at the flanks exposes quartzites and conglomerates of the Seljord Quartzite (Seljord Group), schists of the Bondal Schist (Seljord Group) and calcareous schists of the Lauvhovd Schist (Seljord Group). Furthermore, the rocks of the Rjukan and Seljord Groups have been intruded by thick mafic sills and dykes which have been folded as well. Finally, minor granite intrusions occur in the area. A geological sketch map of the area is shown in Fig.3.

A detailed description of the geology of the Bleka Fold area is given by Svinndal (1952).

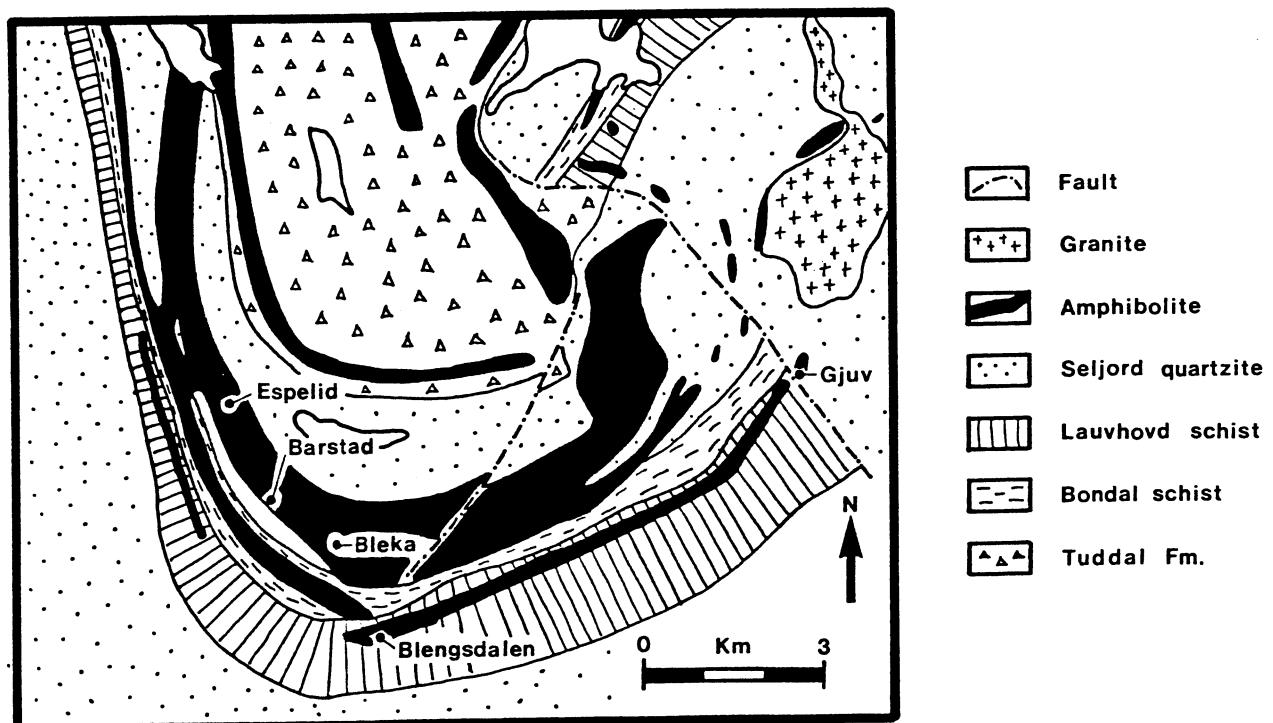


Fig.3. Geological sketch map of the Bleka Fold area. Map mainly bases on Neumann and Dons' map at scale 1:100,000 and on Svinndal (1952).

3. RESULTS.

The investigations performed comprise in addition to a photogeological interpretation detailed field work in selected areas as well as reconnaissance exploration throughout the Bleka Fold area. The work was initially concentrated on investigations in the Bleka Mine area in order to achieve a better understanding of the gold-bismuth bearing vein system, but was later concentrated in the Espelid area due to the finding of a significant vein swarm. The work is described in the following order:

- 1) Lineament analysis of the Bleka Fold area.
- 2) The Bleka Mine area.
- 3) The Espelid area.
- 4) Investigations of other known vein occurrences.
- 5) Reconnaissance for Au-bearing veins in unknown areas.

3.1. Lineament analysis of the Bleka Fold area.

Prior to field work a photogeological interpretation of the area was carried out. The basis of this interpretation is a series of air photos at scale 1:30,000 from Norsk Luftfoto og Fjernmåling I/S (H76 1642 - Lines A-F).

A photomosaic was constructed and two types of lineaments - major distinct lineaments and unspecified lineaments - were interpreted. Furthermore, a few circular features were outlined. In addition to the lineament/circular feature interpretation the amphibolite sills and dykes, which host the Au-Bi veins, were provisionally outlined. This was possible in particular in the western part of the area where a lithologically based vegetation difference is pronounced.

The lineament interpretation is biased in the way that more subtle lineaments have been included from amphibolitic areas. This is due to a hope that even the most insignificant lineament could be a guide to a vein system.

The result of the interpretation is shown in the enclosed Map 1 (scale ~1:20,000) and in Fig. 4 (scale ~1:80,000).

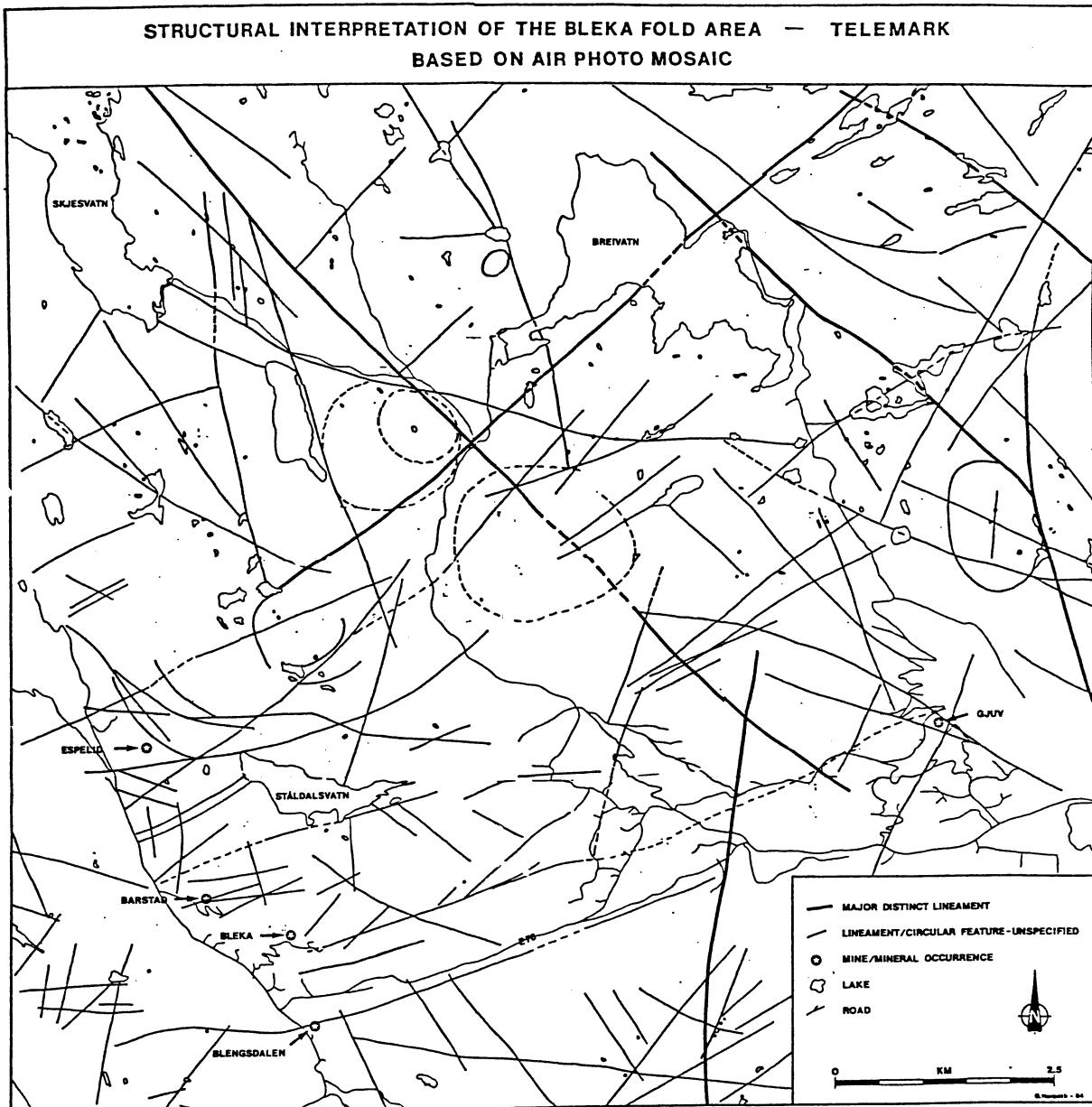


Fig.4.

It is seen that the major distinct lineaments are represented by three main directions:

- a) appr. NW-SE
- b) appr. NE-SW
- c) appr. N-S

Including all interpreted lineaments the two directions - NW-SE and NE-SW are the dominating ones with minor directions NNW-SSE, NNE-SSW, E-W and N-S. This is in accordance with mapped faults of the Telemark area (Sigmund et al., 1984) where in particular the NW-SE faults are prominent structural features. Movement along the three prominent directions seems to be sinistral for the NW-SE and N-S faults and dextral for the NE-SW faults. The N-S faults are clearly the oldest features as they are cut by both the NW-SE and the NE-SW fractures. The NW-SE/NE-SW faults could be interpreted as a conjugate fracture system resulting from E-W compression contemporaneous with the N-S folding of the central Telemark area. If this is to hold true it means that the level in the crust in which the fracturing/shearing is observed was located somewhere close to the boundary zone between the ductile and the brittle regime since the angle between the two fracture directions in the plane of max stress is 60° - 90° (Ramsay, 1980). This means that the gold-bismuth bearing Bleka-type veins which strike 40° - 90° could be related to dextral shear faults associated with E-W to ESE-WNW compression. In fact, as will be demonstrated later this is the actual case.

However none of the more pronounced interpreted lineaments (40° - 90°) within amphibolitic terrain which were investigated in the field could be directly related to vein systems. This is probably mainly due to the very low overall exposure (5%), but could as well be due to lack of relation as implied by the known and newly discovered veins which do not show up as lineaments on air photos.

In summary, it can be concluded that there is no direct relationship between interpreted lineaments and observed vein systems. However, the lineament analysis is a powerful means of unraveling the structural evolution of the area and as such an indirect guide to vein mineralization.

3.2. The Bleka Mine area.

The Bleka Mine area as defined by this investigation is outlined in the enclosed Map 2. It covers an area of approximately $\frac{1}{2}$ km² and includes the 1100 m known strike length of the Bleka Main Vein System as well as associated veins.

The area has been covered by both surface and subsurface investigations as outlined below.

The history of the mine and the previous work carried out is reported by Vogt (1886, 1888), Kostke (1902), Wielgolaski (1918), Bugge (1935), Horvarth (1943), Hake (1943a, 1943b), Lauer (1943), Svinndal (1952) and Pedersen (1984). The work by Vogt and by Kostke is by far the most comprehensive and in particular the early work of Vogt gives an excellent description of the geology of the mineralized veins. A summary of the most important reported results is presented in Table 1.

The surface investigations carried out include mapping and detailed sampling of the Bleka Main Vein System along the known 1100 m strike length as well as mapping and detailed chip sampling of associated veins. The main results of this are presented in the enclosed Map 2. The area has been mapped at a scale of 1:1000 based on enlargement of 1:5000 maps from Økonomisk Kartverk. As the overall exposure of the amphibolite is as low as 5-10% probably only a small portion of the existing veins and veinlets have been located in the area (Vogt (1888) notes that appr. 20 veins were known in the Bleka area). The trace of the Bleka Main Vein is only mappable due to old trenching of the vein. Most of the other veins indicated on the map were also located due to trenching and blasting. As it is seen in Map 2 the Bleka Main Vein is only exposed in a few localities (less than 10) otherwise covered by filled-in material along the old trench or not exposed. The vein is estimated to exhibit an average surface thickness in the range of 20-25 cm over a total length of more than one km. From the map it is obvious that a major shift in orientation occurs along a 300 m long part centered around the old workings. This is probably due to the convergence of two en echelon fractures. This point is further elaborated in section 4. The mean orientation of the vein is 71°/73°N (43 measurements) which is in accordance with earlier results. Macroscopically, the main vein is a quartz-tourmaline-ankerite vein with minor calcite, dolomite, epidote, muscovite and chlorite. The ore minerals which in general constitute around 1%, but occasionally up

YEARS	ACTIVITIES	IMPORTANT NOTES AND RESULTS	REFERENCES
1880-1900	Mining	Small scale mining by Compagnie Francaise des Mines de Bamble. No detailed report exists on the results of the mining activity. Excellent geological description of the veins.	Dons (1963) Vogt (1886,1888)
1900-1935	Sporadic geological investigations and feasibility studies.	Geology: The Bleka Main Vein is an irregular vein, which has been followed along 1100 m's length. The average thickness - 35 cm and the average orientation 74°/70°N. The vein often exhibits a banded structure with tourmaline towards the margins and carbonates in the centre. Late often cross-cutting quartz veinlets carry bismuthinite and visible gold. There is clearly an intimate relationship between bismuth and gold. Mine data: In 1901 a total of 701 m of drifts existed. An estimate of the vein tonnage above mine level 525 amount to appr. 50,000 t. A very thorough sampling programme carried out in the mine in 1901 totalling appr. 1800 kg of vein material was split into several representative samples totalling 150-200 kg yields an average grade of the undiluted vein of 25 ppm Au, 43 ppm Ag, 0.8% Bi (a more realistic value according to the analysis results is 0.3-0.4% - see Kostke) and 1.4% Cu. This average estimate is calculated on a tonnage of 16,000 t of ore. Older analyses (1880-1890) show the high grade nature of the vein as selected samples yield values in the range 50-700 ppm Au (one sample - 5.3% - possibly a sort of concentrate) and 100-2000 ppm Ag.	Kostke (1902)
		A feasibility study in 1918 does not include any important new information.	Wielgolaski (1918)
		In a publication by Bugge an assay of a 12 kg sample shows 14.5 ppm Au, 30 ppm Ag, 0.42% Bi, 0.23% Pb and 275 ppm Cu. Bugge suggests that galenobismuthite is the main bismuth mineral of the ore.	Bugge (1935)
1935-1939	Mining.	Mining was carried out by A/S Bleka Gruber. No reports exist on this activity. According to Halvor Johnsen Svarddal (personal communication, 1984) 40 t of ore concentrate (from jiggling) were sent to Germany during this period. A shaft was made 70 m down from mine level 525 (drift F of Kostke (1902)) and three lower levels were investigated. According to Halvor Johnsen Svarddal very low gold values were encountered.	Dons (1963)
1940-present	Feasibility studies and geological investigations	During the Second World War the Germans showed renewed interest in the Bi-potential of the Bleka Mine. A report by Horvarth includes a very optimistic ore estimate of 60,000 t with a grade of 0.7% Bi, 0.6% Cu, 11 ppm Au and 40 ppm Ag. Two reports by Lauer and Hake downgrades this estimate drastically. Lauers geological investigations of the underground workings conclude the following: the gold-bismuth mineralized lenses (mainly the thicker parts of the veins) within the Bleka Main Vein in general have a strike length of 20-50 m with thin barren parts inbetween. They are mainly composed of tourmaline and ankerite with only minor amounts of quartz. From the description it seems as if there is a distinct zonation with predominantly quartz in the lower and western parts of the mine and predominantly tourmaline-ankerite in the upper and eastern parts. The major stoped-out lens of level 560-580 (drift B and A) and the stoped-out lens of level 525 (drift F) are not interconnected and thus represent individual lenses. The inner part of the four major drifts (A-B-C-F) are all very low grade with respect to gold and bismuth. Of the 615 m drift which could be investigated, 120 m were stoped out (the main ore lenses), 307 m were unmineralized (mainly developed as a very thin veinlet or as a fault plane) and 190 m represented a true, but thin vein without any visible gold and bismuth mineralization. This last part representing approximately 200 m of weakly mineralized vein was thoroughly sampled and split into 37 representative samples (each representing from 3-10 m strike of vein). A total of 2.5 t of vein material was sampled. The samples were crushed and split into 2-3 kg samples which were analyzed at Kongsberg Silvermines for Au, Bi and Cu. The average grade of this left-over part of the vein system is 1 ppm Au, 400 ppm Bi and 0.25% Cu. The individual analyses and a sketch of the mine with indicated sample points are shown by Pedersen (1984 Fig.58). On the basis of this discouraging result no further recommendations are made for the Bleka Mine by the Germans.	Horvarth (1943) Lauer (1943) Hake (1943)
		The work by Svinndal adds no new information on the Bleka Mine, but mentions that pure tourmaline veinlets are found in the area.	Svinndal (1952)
		Pedersen visited the mine on a reconnaissance basis and stresses the potential of the mineralization because of the steep dip of both the veins and the amphibolite. Furthermore, prominent stream-sediment anomalies of gold were outlined in the Bleka Fold area.	Pedersen (1984)

Table 1. Summary of the history of the Bleka Mine.

to several per cent, comprise chalcopyrite and pyrite with subordinate bismuthinite, Bi-sulfosalts, gold, galena and scheelite. The economically interesting minerals gold, bismuthinite and Bi-sulfosalts are in particular concentrated in the eastern part of the vein, but the Bi-minerals are also quite widespread in the Sverveli area in the westernmost part (Map 2). Visible gold was only observed in two samples (2110061 and 2110062) both originating in the central mine area around drifts 560 and 580 (B + A according to Kostke (1902)). The gold grains observed are up to 1 mm in the greatest dimension and occur partly in late quartz veinlets and partly intergrown with tourmalin-quartz, in both cases associated with bismuthinite. Galena was only observed in the Sverveli area where it occurs in both the main vein and in adjoining feather veins. The Sverveli area is situated at a level approximately 100 m higher than the main ore lenses of the vein system and thus could represent a lower temperature ore assembly.

Pronounced hydrothermal alteration is associated with the veins. This alteration has been thoroughly investigated by Vogt (1888) who describes it as a zone of appr. 0.5 m's thickness on either side of the veins. Mineralogically, the amphibolite host rock increases its muscovite and calcite content towards the vein, whereas the hornblende and feldspar content decreases. This results in an overall bleaching of the rock towards the vein. Eventually, the altered amphibolite is transformed into a quartz-muscovite-calcite rock. Furthermore, development of idiomorphic pyrite cubes and magnitite octahedrons as well as occasional tourmalinization and ankeritization is a characteristic feature. Minor chalcopyrite impregnation is locally developed. This progressive hydrothermal alteration process is almost identical with alteration phenomena described by Robert and Brown (1984) who investigated gold-bearing quartz-tourmaline veins at the Sigma Mine in Canada. Figs. 5-7 exemplify the progressive alteration observed at the Sigma Mine. The processes observed are virtually identical to those identified at Bleka. However, an important difference seems to be the distribution of gold within the altered zones. At Sigma, gold enrichment is associated with intense carbonatization in the zone of visible alteration. At Bleka this has not yet been observed,

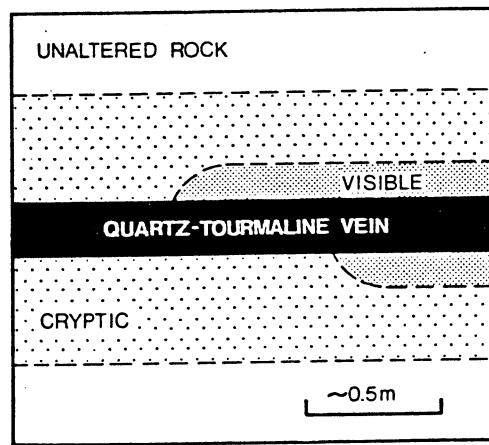


Fig.5. Schematic representation of distribution of cryptic and visible alteration zones around veins and stringers at the Sigma Mine. Visible alteration may be absent and grade laterally to asymmetrical or to symmetrical alteration envelopes. From Robert and Brown (1984).

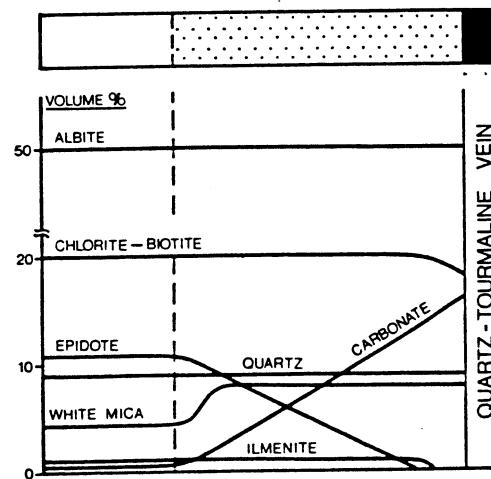


Fig.6. Idealized semiquantitative distribution of minerals involved in cryptic alteration. From Robert and Brown (1984).

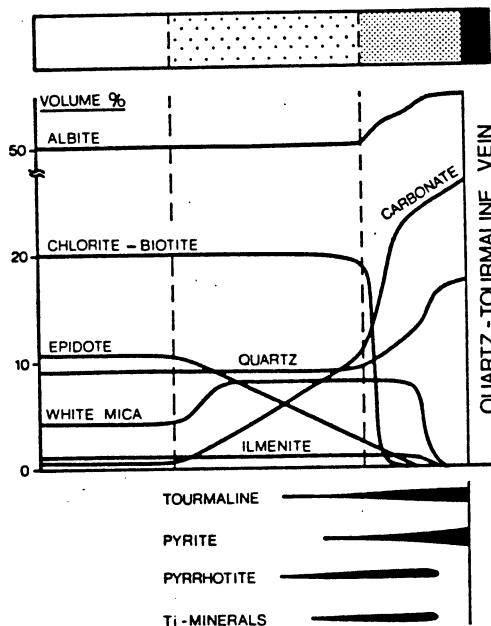


Fig. 7. Idealized semiquantitative distribution of minerals involved in visible alteration superimposed on cryptic alteration. From Robert and Brown (1984).

although several samples of the alteration zones have been analyzed for gold. Fig. 8 shows visible alteration around a small Bleka-type vein.

The macroscopic appearance of the exposed main vein is shown in Figs. 9-12.

The other veins found in the Bleka Mine area are probably all associated with the fracture system of the Bleka Main Vein. Most of these veins occur within a short lateral distance (<100 m) from the main vein and are relatively thin (2-20 cm). In only one of the veins visibly Bi-mineralization was observed (sample 2110046 - Map 2) whereas the rest mainly represent quartz-tourmaline veins with minor pyrite and chalcopyrite.



Fig.8. Visible alteration (bleaching) around a small Bleka-type vein. Locality close to sample 2110046 (Map 2).



Fig.9. Outcropping 10-20 cm thick Bleka Main Vein at the entrance to mine level 530.



Fig.10. Outcropping 20-30 cm thick Bleka Main Vein close to the shaft level 595. Contains locally abundant Bi-Au mineralization.



Fig.11. 30 cm thick Bleka Main Vein at the entrance to mine level 560. Mainly composed of ankerite and tourmaline.



Fig.12. 10 cm thick vein of the Bleka Main Vein System close to mine entrance 550. Swelling of the vein due to dextral shear movement is evident.

The subsurface investigations carried out include inspection of approximately 300 m of accessible drifts and sampling of selected vein sections. Only mine levels 610, 560, 550 and 525 were accessible for inspection. The results of the inspection are in good agreement with those of Kostke (1902) and Lauer (1943) (see Table 1). As could be expounded from Lauers descriptions there is a distinct zonation with predominantly quartz in the lower and western parts of the mine and predominantly tourmaline-ankerite in the upper and eastern parts. Figs.13-15 exemplify this. In the upper mine levels the vein often exhibits a banded appearance with tourmaline rimming ankerite (Fig.15). No visible gold mineralization and only very minor visible bismuth mineralization was observed in the mine. However, the stoped areas were not accessible thus only the poorly developed and weakly mineralized vein sections could be investigated.



Fig.13. One metre thick quartz-tourmaline vein from stoped area of level 525 (drift F). Note internal Riedel shear planes with tourmaline and sulphides. The shear planes are a result of dextral shear movement and the seeming sinistral appearance is because the photo is taken of a roof exposure.

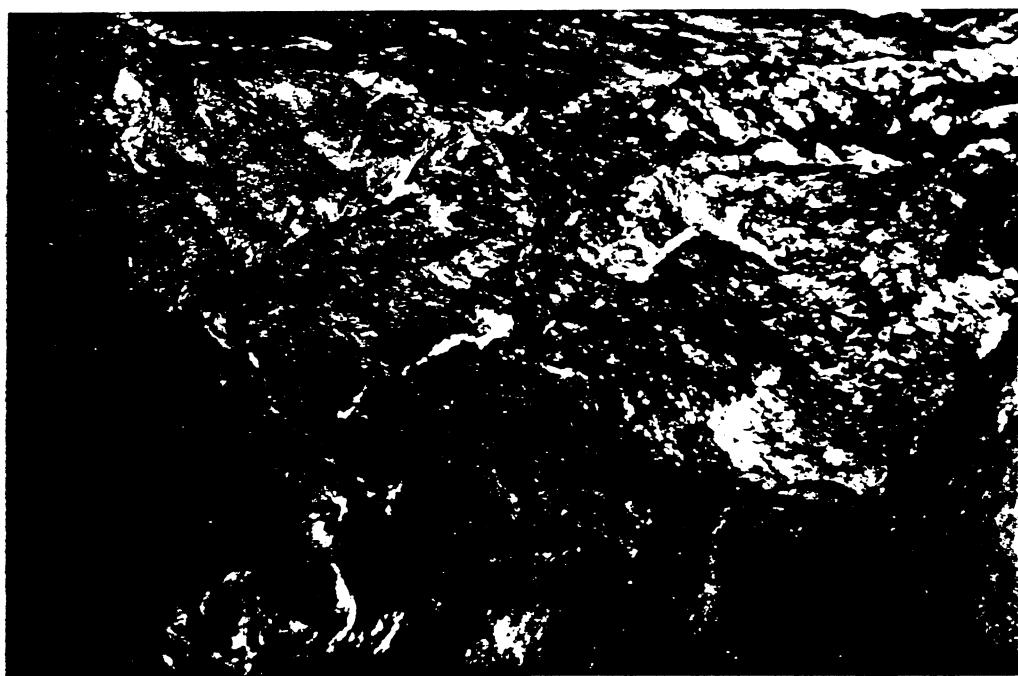


Fig.14. 30-50 cm thick tourmaline-ankerite vein with late quartz veinlets. From the upper mine level 610 (drift C)

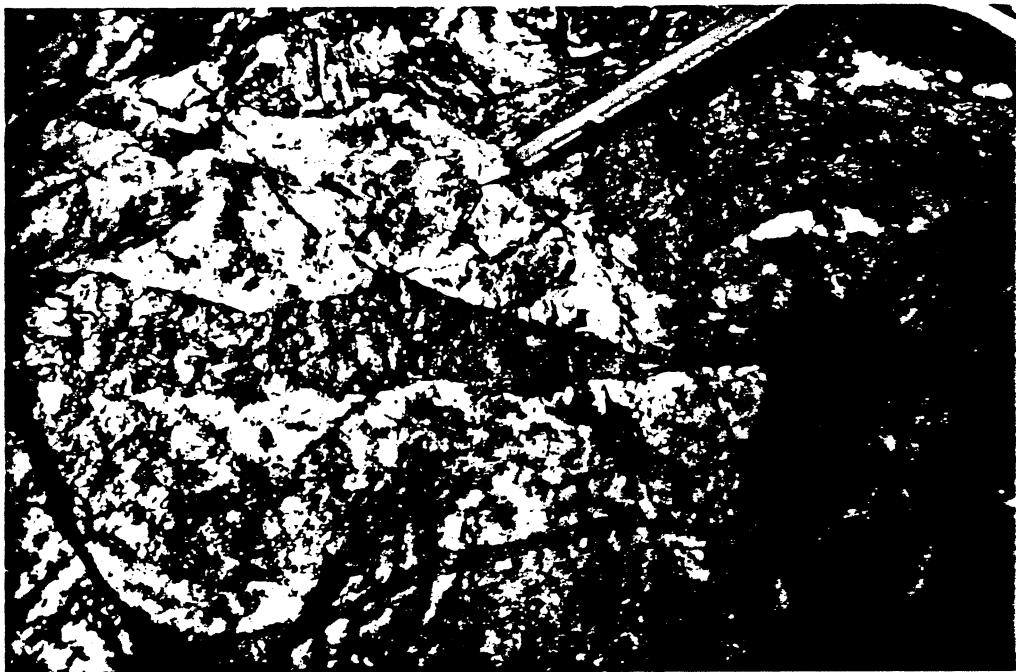


Fig.15. Quartz-ankerite-tourmaline vein from mine level 560 (drift B). Note rimming tourmaline around ankerite veinlets.

Structurally, the vein is complex. It pinches and swells over short distances and locally disappears laterally into a 5-10 cm mylonitic fault plane with abundant chlorite and calcite. The shear movement associated with the vein emplacement is clearly dextral (Figs.16-17) and a detailed study of the geometry of the feather joints demonstrates that shear faulting took place within the brittle-ductile shear regime (Fig.17).

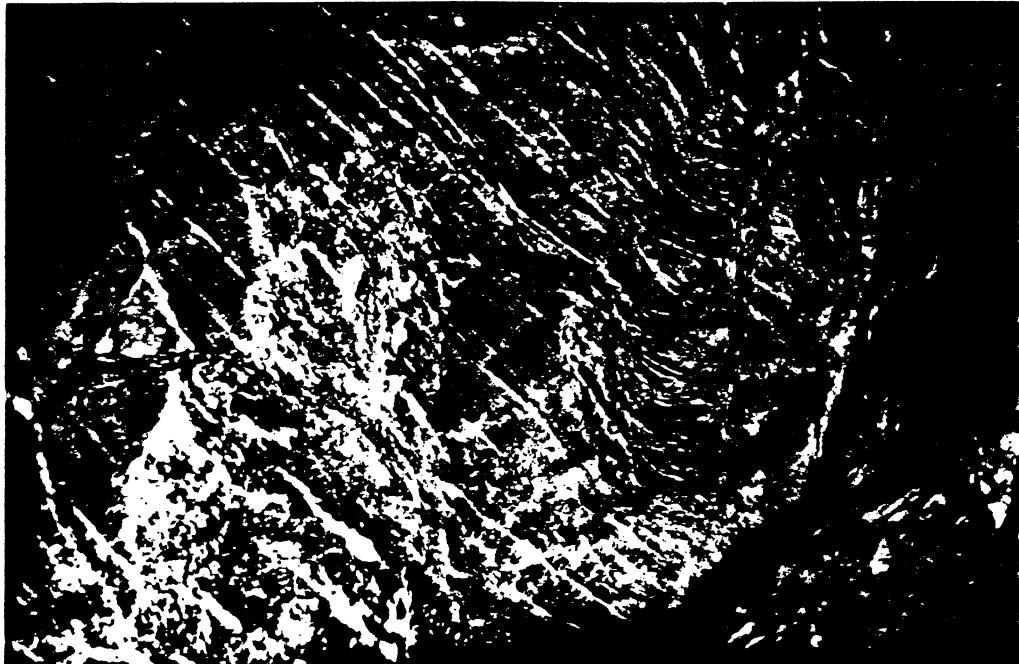


Fig.16. Major system of pinnate calcite veinlets associated with the Bleka Main Vein. The measured strike of the Bleka vein is 50° and the strike of the joints is 85° - indicative of dextral shear. From mine level 560 (drift B). Photo from roof exposure.

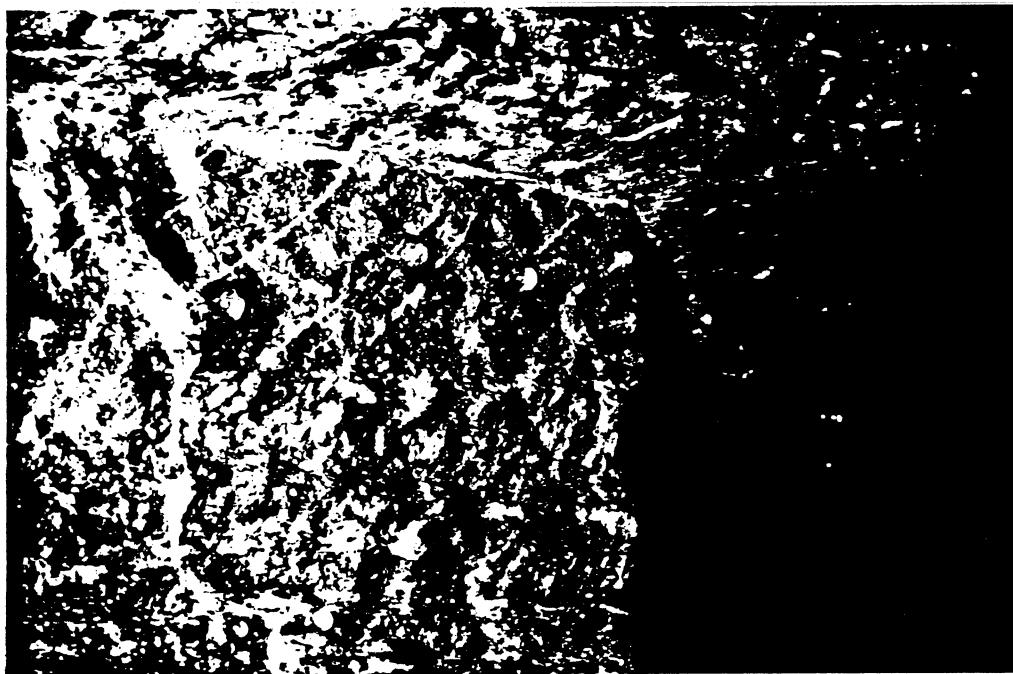


Fig.17. Minor system of calcite veinlets associated with the Bleka Main Vein. Exemplifies dextral shear within a brittle-ductile shear regime. The seeming sinistral appearance is because the photo is taken of a roof exposure. From mine level 560 (drift B).

Ore microscopy. Only two polished sections have been prepared and investigated. Prior to the field work a section was prepared from litho-sample 020107A (Pedersen, 1984) with a gold content of 72.7 ppm (average of four analyses). The primary ore minerals identified include chalcopyrite and bismuthinite (both intergrown with cpy and as isolated grains) and subordinate pyrite, sphalerite, rutile and native gold. Only two small (10-20 μ) grains of gold were found isolated in micro-fractures in quartz. The observed amount of gold only corresponds to a gold content of 1-2 ppm and clearly illustrates the nugget effect of the Bleka vein mineralization even on a small scale.

A polished section of a highgraded pan concentrate of crushed ore from the jigging table of the old milling plant has also been prepared. The ore minerals identified include the major constituents native gold, bismuthinite and native bismuth and subordinate chalcopyrite, pyrite, Bi-sulfosalts (cosalite?, aikinite?, galenobismuthite?, emblectite?), maldonite?, bornite and galena? (probably galena with a high Bi-Ag content). In addition, a UV-light examination showed a high content of scheelite.

Native gold occurs mostly as anhedral grains and aggregates (Figs.18-19), but subhedral and euhedral crystals occur as well (Figs.20-23). The observed grain size (of single grains or aggregates) varies considerably from <10 μ up to 2.5 mm. The luminous golden yellow colour of the gold also varies considerably (Figs.18, 19 and 27) indicative of a great variation in the silver content. A few more reddish gold grains could be indicative for a high bismuth content. Gold is observed intergrown with native bismuth (Figs.20-23), with bismuthinite and/or Bi-sulfosalts (Figs.24-26), with chalcopyrite (Fig.26), with maldonite? (Fig.21), with Bi-bearing galena? and with tourmaline (Fig.27). A genetically very interesting^{feature} is the common coarse myrmekitic intergrowth between gold and bismuth (Figs.20-21). The probable occurrence of maldonite (Au_2Bi) associated with the myrmekitic intergrowth is limited to a narrow margin within the high temperature range (Ramdohr, 1980). This texture is observed at the classical occurrence of maldonite (Maldon Mine, Australia) which is a very high temperature gold-quartz vein.

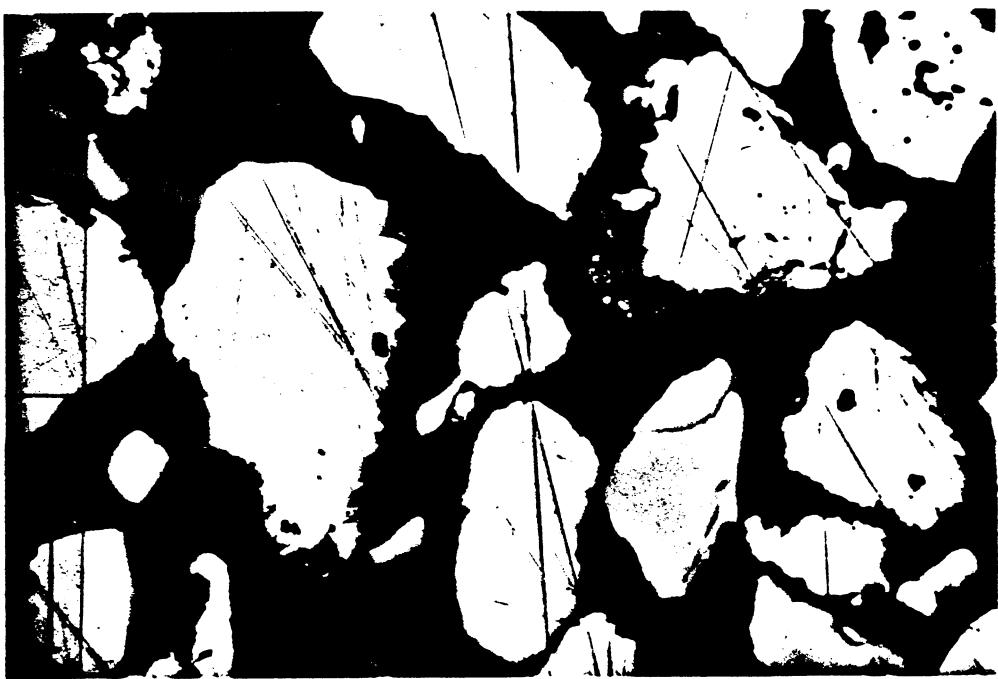


Fig.18. Photomicrograph showing anhedral gold grains. Varying Ag-content is reflected by the colour differences. Oil immersion. Bar scale 0.1 mm.



Fig.19. Photomicrograph showing anhedral-subhedral gold grains. Varying Ag-content is reflected by the colour differences. Native bismuth and bismuthinite occur as well. Oil immersion. Bar scale 0.1 mm.



Fig.20. Photomicrograph showing myrmekitic intergrowth between native gold and native bismuth. Oil immersion. Largest dimension of grain - 0.3 mm.



Fig.21. Photomicrograph showing myrmekitic intergrowth between native gold, native bismuth and malodinite? (bluish). Oil immersion. Largest dimension of grain - 0.2 mm.

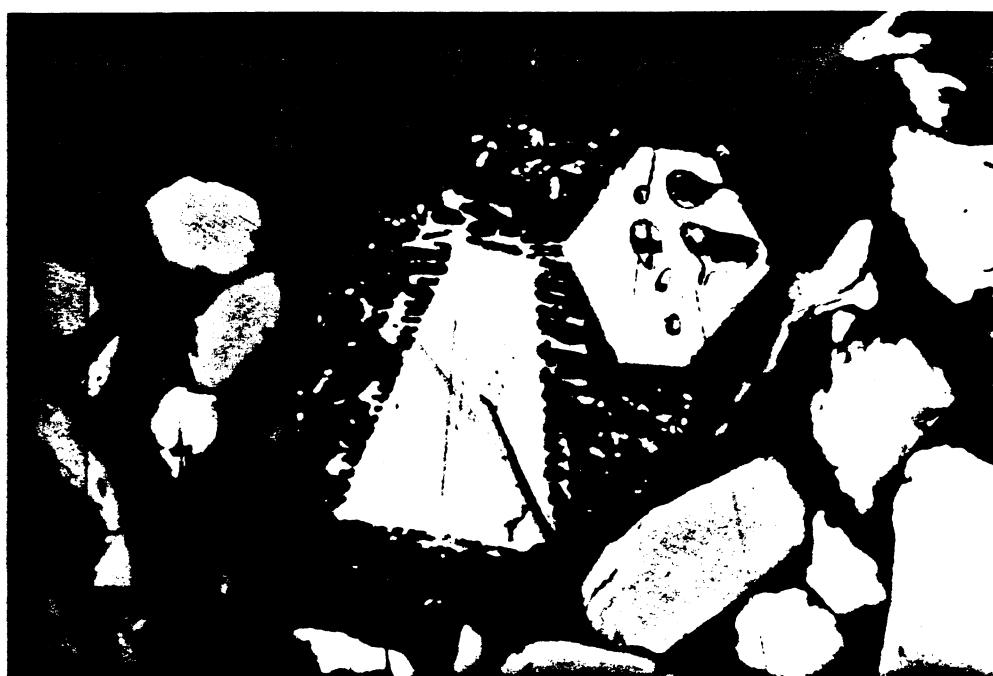


Fig.22. Photomicrograph showing aggregate with euhedral gold dodecahedron and native bismuth crystal with "feathers". Oil immersion. Largest dimension of gold crystal - 0.1 mm.

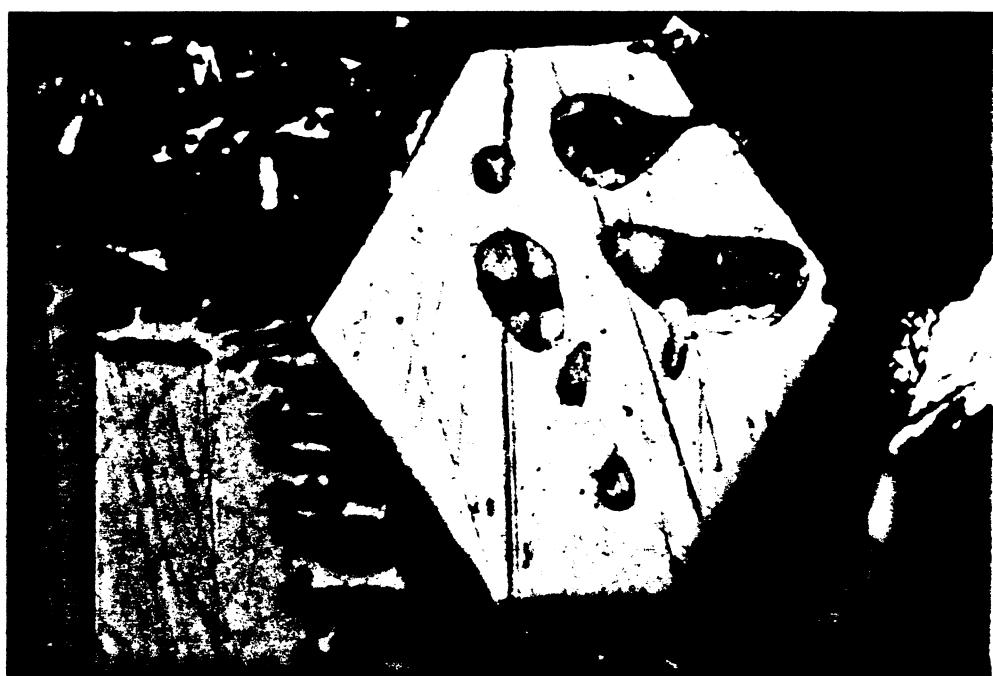


Fig.23. Enlargement of part of Fig.22. Euhedral gold dodecahedron with trapped bismuth droplets. Oil immersion. Largest dimension of gold crystal - 0.1 mm.

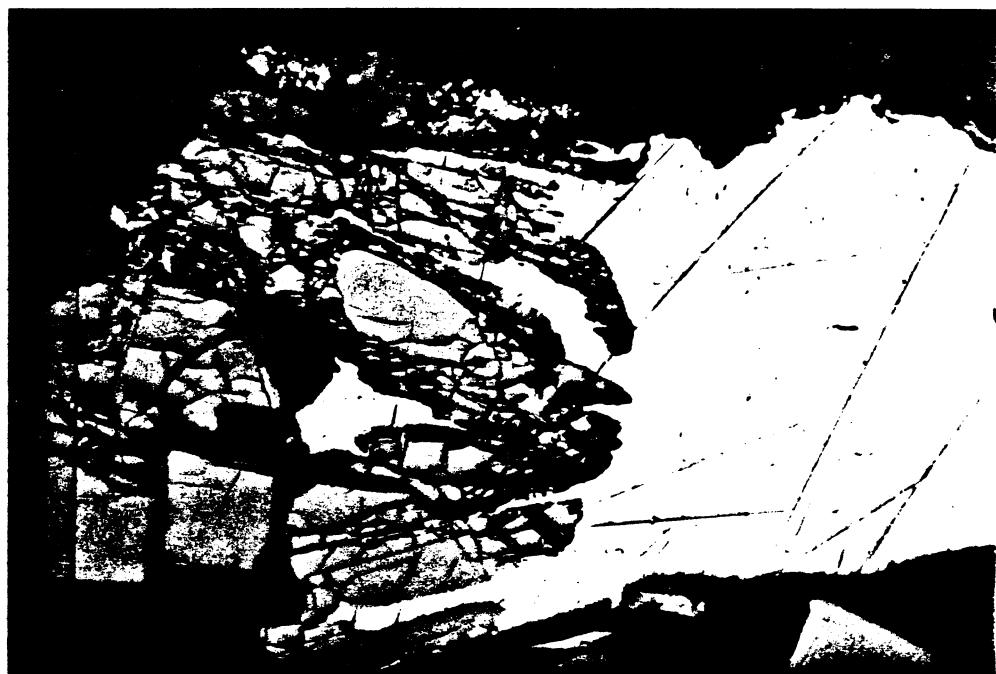


Fig.24. Photomicrograph showing intergrown gold and bismuthinite. Oil immersion. Size of gold grain 0.2·0.5 mm.



Fig.25. Photomicrograph showing intergrown gold, bismuthinite and Bi-sulfosalts. Oil immersion. Size of gold grain 0.1·0.5 mm.

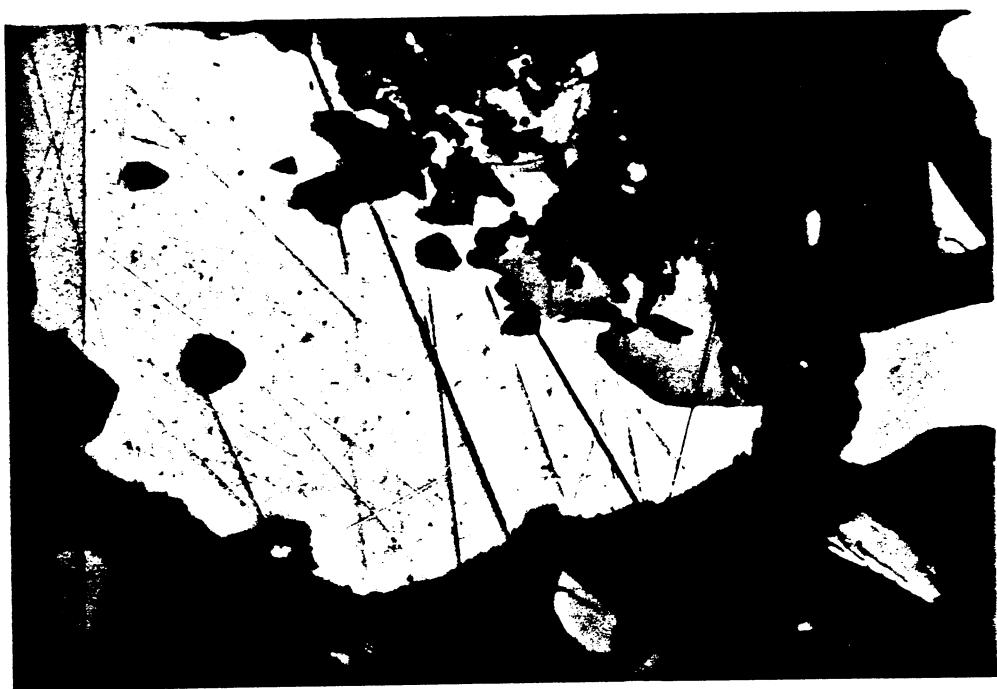


Fig.26. Photomicrograph showing intergrown gold, bismuthinite and chalcopyrite. Oil immersion.
Size of gold grain 0.2-0.5 mm.



Fig.27. Photomicrograph showing gold with inclusions of tourmaline. Note varying Ag-content of the two major gold grains. Oil immersion. Size of major gold grain 0.2-0.4 mm.

The occurrence of native bismuth and maldonite (?) and the intimate myrmekitic intergrowth with gold has not been reported elsewhere. However, the intimate relationship between gold and bismuth minerals reported as early as 1886 is further supported by this preliminary microscopic investigation.

Analysis results.

In summary, the old analyses indicate the following undiluted grades:

High grade ore lenses (- 10,000 t) :

Au - 25 ppm
Ag - 43 ppm
Bi - 0.4 %
Cu - 1.4 %

Low grade vein mineralization (- 15000 t) :

Au - 1 ppm
Ag - -
Bi - 400 ppm
Cu - 0.25 %

Geochemically, this means that the Bleka Vein is a bismuth-gold vein as indicated by the following enrichment factors for the high-grade lenses:

Enrichment factor	Element
500,000	Bi
10,000	Au
500	Ag-Cu
10-100	W-Pb-Zn

The analysis results of grab and chip samples from both surface and subsurface investigations are presented in Appendix II and the gold analyses are plotted in the enclosed Map 2. Analysis of selected samples from the Bleka Main Vein confirms the mapping results. Only samples from the central mine area have high gold contents (up to 94 ppm) whereas samples from outside this area only show gold values

in the range of < 20 ppb to 1.7 ppm. Besides gold, interesting Ag-values are encountered in sample 2110061 (> 250 ppm) of the mine area, but in particular the western (and higher) area around Sverveli exhibits interesting silver values (100- > 250 ppm). Silver is clearly associated with high lead and bismuth values and is probably contained within galena (in particular in the Sverveli area) and in Bi-sulfosalts (e.g. sample 2110001).

In general, bismuth seems to be a much more efficient element in delineating mineralized vein systems than gold. Most of the veins analyzed for both gold and bismuth in the Bleka area show distinct bismuth enrichment (10-100 ppm) without any detectable gold (< 10-20 ppb). This means that the bismuth halo within the vein system is much more pronounced and extensive than that of gold. Thus, Bleka-type veins of the fold area should preferably be analyzed for both gold and bismuth and any encountered bismuth anomalies without gold would probably be indicative for a close gold-bismuth mineralization.

3.3. The Espelid area.

The Espelid area is outlined in Fig.1 and in the enclosed Map 3. It covers an area of approximately 3/4 km² from the Espelid Farm in the NW to just north of Livesengåsen in the SE and is situated from 2-3½ km NNW of the Bleka Mine. Geologically, the setting is similar to that of the Bleka area except for one important difference - the Bleka area is situated close to the axis of the fold whereas the Espelid area is located at the western limb of the fold. This means that the 40°-90° striking veins of the area are virtually perpendicular to the strike of the amphibolite sill (Fig.28). Furthermore, the thickness of the sill is only half of that observed in the Bleka area where it reaches a max thickness around 1000 m. This may partly be a result of primary variation in thickness and partly a result of boudinage which may also have effected the degree of dilation in the Espelid area (Fig.28). However, the steep dip of both the veins and the amphibolite sill makes room for a significant tonnage at depth.

The area was investigated on a reconnaissance basis partly due to an old (1880) Cu-bearing quartz vein occurrence (Espelid) and partly due to stream sediment anomalies of Au-Bi-Pb-Zn-Co reported by Pedersen (1984).

The old trenching which seemed to have been left untouched for a great many years occurs approximately 500 m SE of the Espelid Farm and only 25 m N of the Nystaul saeter cottage. The occurrence is exemplified by samples 2110033-36 on the enclosed Map 3. The Espelid veins seem to be part of a major vein swarm identified during reconnaissance exploration of the area. Due to this encouraging find the area was mapped at scale 1:1000 (based on enlargement of 1:5000 maps from Økonomisk Kartverk) and the results of this are presented in the enclosed Map 3. As it is seen in Fig.28 and in Map 3 an area of $1000 \cdot 500 = \frac{1}{2} \text{ km}^2$ contains in the order of 10 veins with locally more than $\frac{1}{2} \text{ m}$'s thickness (max observed thickness = 2 m - Fig.29), 20 veins with a max thickness of 10-50 cm and a greater number of veins with max thickness of 1-10 cm. The veins are assumed to have a strike length of 400-500 m which is the approximate thickness of the sill. One of the major veins (samples 2110116-133, Map 3 - Fig.30) could be followed more or less continuously over 250 m laterally and 75 m vertically with an average thickness of 75 cm. Due to the very low exposure ($\sim 5\%$) of the amphibolite the estimated amount of veins and the length of these is very uncertain, but it is assumed that additional medium sized veins occur in the area. In general, the size of the hydrothermal vein system indicates an overall dilation much greater than observed elsewhere in the Bleka Fold area.

The veins are very similar to those of the Bleka mine area. The average orientation of the veins is $68^\circ/78^\circ \text{N}$ (average of 141 measurements - standard deviation of strike = 15°) with most of the veins striking $40^\circ-90^\circ$. The veins consist predominantly of quartz and tourmaline with varying amounts of ankerite-sericite-pyrite and locally chalcopyrite (Figs.31-32). In general, the overall sulphide content is lower than in the Bleka Main Vein. Pronounced hydrothermal alteration similar to that described from Bleka is ubiquitous.

Characteristically, evidence of dextral shear movements are found throughout the vein swarm. As seen in Fig.34 much of the observed shear movement is lateral due to the frequent

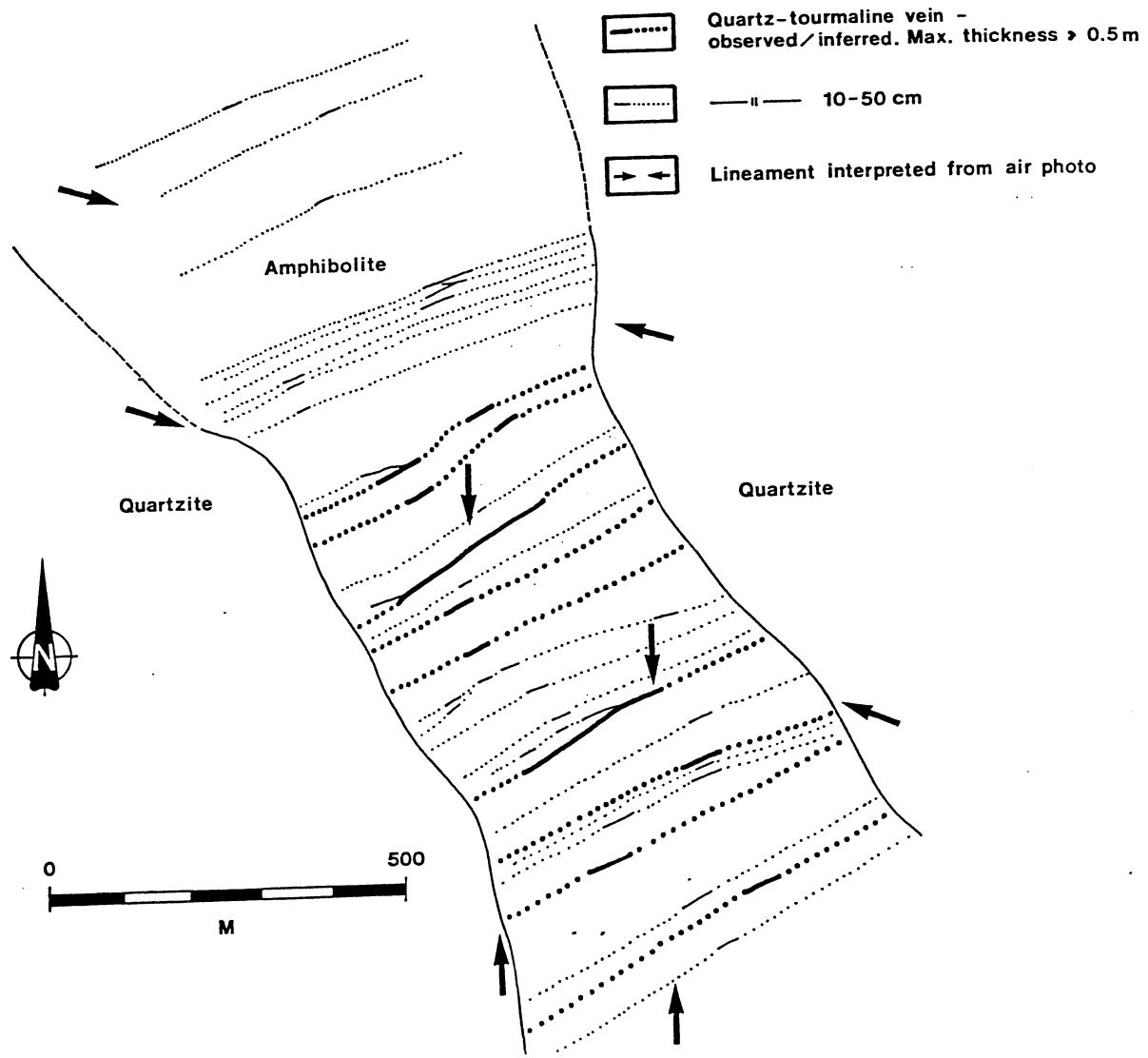


Fig.28. Simplified sketch of the Espelid vein swarm.

occurrence of horizontal slickensides. The dextral nature of the shear movement is seen as widespread feather or pinnate veins (Figs.35-37).

Due to the great similarity with the veins of the Bleka Mine area a detailed chip sampling programme was carried out covering almost all outcrops of veins with a thickness greater than five cm. The locations of the sampled outcrops are all shown on Map 3. More than 110 chip samples (1-10 kg according to vein thickness) and supplementary grab samples were collected. All of these samples have been analyzed for gold and



Fig.29. Up to 2 m thick
quartz-tourmali-
ne vein of the
Espelid Vein
Swarm. Sample
2110231 - Map 3



Fig.30. 1½ m thick quartz-
tourmaline vein
of the Espelid
Vein Swarm. Sam-
ple 2110121 -
Map 3.



Fig.31. Close-up of 30 cm thick quartz-tourmaline vein with abundant pyrite. From the Espelid Vein Swarm. Sample 2110105 - Map 3.



Fig.32. Close-up of a 10 cm thick quartz tourmaline-ankerite vein with pyrite-chalcopyrite and probably Bisulfosalts. From the Espelid Vein Swarm. Samples 2110141 and 2110201-202. Note wrong sample number!

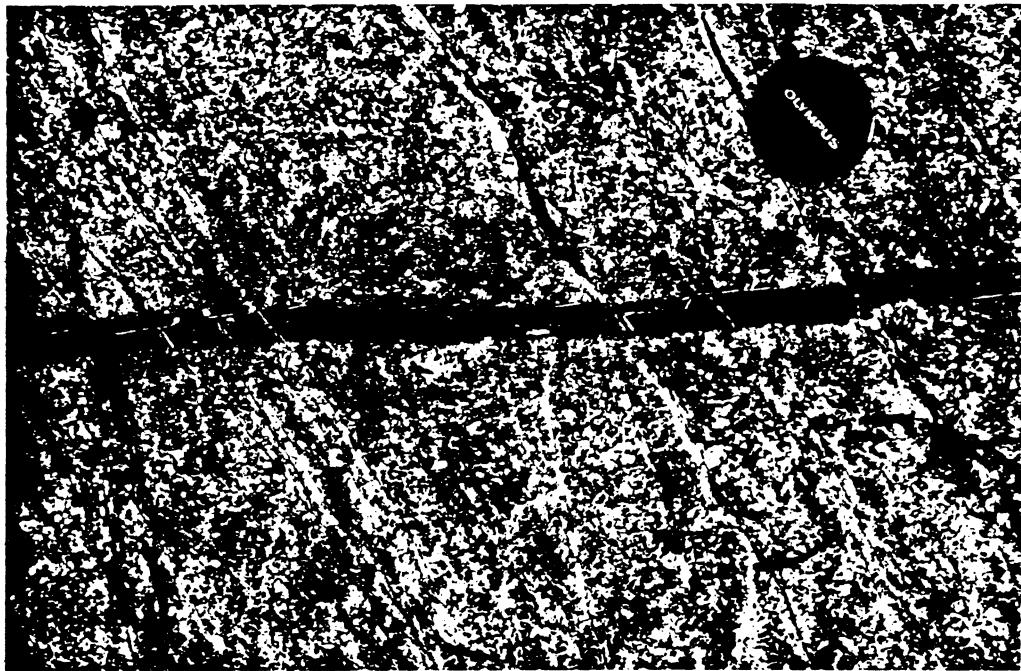


Fig.33. Thin tourmaline vein in amphibolite in the northern part of the Espelid Vein Swarm where these are quite widespread.



Fig.34. $\frac{1}{2}$ m thick quartz-tourmaline vein with pronounced horizontal slickensides. From the southern part of the Espelid Vein Swarm. Sample 2110220 - Map 3.



Fig.35. 10 cm thick quartz-tourmaline vein with adjoining feather veins indicative of dextral shear. From the Nystaul occurrence. Sample 2110038 - Map 3.



Fig.36. 80 cm thick quartz tourmaline vein with feather veins indicative of dextral shear. From the Espelid Vein Swarm. Sample 2110228 - Map 3.



Fig.37. 25 cm thick quartz-tourmaline vein with feather veins indicative of dextral shear. From the Es-pelid Vein Swarm. Samples 2110237-238 - Map 3.

approximately 20 samples for the additional elements Bi-Ag-Cu-Pb.

The analysis results indicate that a Au-Bi anomalous area covering $\frac{1}{2} \cdot \frac{1}{2} = 1/4 \text{ km}^2$ occurs around Nystaul-Reiustaulhaugen (Map 3). Although the values are low compared with the contents of the Bleka Main Vein a distinct enrichment in both gold (20-210 ppb) and bismuth (2-144 ppm) is observed. Samples from outside this area are low in both elements. Also lead seems to be relatively enriched, whereas copper does not show any particular enrichment in this area.

On the basis of these results it is suggested that the gold-bismuth enrichment represents a halo-phenomena to a deeper-lying gold-bismuth mineralization. The low Au-Bi contents observed in the SE part of the vein swarm do not necessarily imply barren vein mineralization, but could as well be a result of even deeper Au-Bi mineralization.

3.4. Investigations of other known vein occurrences.

Barstad.

The location of the Barstad occurrence is shown in Fig.1. More precisely the old drift is located 150 m NW of the farm Barstad øvre close to the stream at 550 masl (samples 2110010-12 - Sample Map 5). The entrance to the drift is seen in Fig 38.



Fig.38. Entrance to the 15 m drift at the Barstad occurrence.

The geological setting is similar to that of Bleka and Espelid.

The drift was investigated all along the 15 m length. The vein which is oriented 90°/60°N is locally up to 60 cm thick, but the average is much lower. It consists of quartz, tourmaline (locally very rich with up to 2 cm long crystals) and ankerite with subordinate pyrite, chalcopyrite and possibly bismuthinite. According to Kostke (1902) later quartz veinlets with bismuthinite and visible gold was observed during drifting. This was not observed neither in the tunnel nor in mineralized vein samples of the old dump. At the end of the drift the vein is displaced by a 145° fault.

Two chip samples and one grab sample were collected at the occurrence. Analyses show gold (up to 30 ppb) and bismuth (up to 12 ppm) enrichment.

Blengsdalen.

The location of the Blengsdalen occurrence is shown in Fig.1. More precisely two different veins (150 m apart) have been trenched and blasted over a composite length of 50-100 m. The southern vein is located close to the Blengsdalen farm and is exposed in the small road between the lower and the upper buildings whereas the northern vein is located 150 m NNW of the main building. According to Vogt (1888) the occurrence was discovered in 1884 and tested in 1896.

The geological setting is slightly different from the Bleka area. The amphibolite sill hosting the veins (the Kasin amphibolite according to Svinndal, 1952) is only 100-200 m thick and has intruded the Lauvhovd limey quartz schists, whereas the Bleka amphibolite has intruded the quartzites. However, the Kasin amphibolite host is just a higher level sill of the Bleka amphibolite (Fig.3).

The northern vein has been trenched and blasted over a length of 50-60 m. Only very limited exposure is seen. The vein is oriented 70°/subvertical S and consists of quartz and ankerite with minor tourmaline, chalcopyrite, pyrite and bismuthinite. Pronounced hydrothermal wall rock alteration occurs. The vein probably pinches and swells over short distances and at the southern end a width of one metre was measured. One bulk chip sample of both visibly mineralized and unmineralized chips

was collected from the old dump. Further, two mineralized grab samples were collected. The analysis results indicate surprisingly high gold contents. The two grab samples have gold values of 103 ppm and 1.8 ppm respectively whereas the bulk chip sample shows 0.31 ppm Au, 84 ppm Bi, 0.28% Cu and 3 ppm Ag. The southern vein has been blasted at several locations over a length of 50-75 m. The vein is very irregular and has an average orientation of 70°/ subvertical. In the road cut the vein has a thickness of 20 cm and consists of quartz, tourmaline and ankerite with pyrite, chalcopyrite and bismuthinite (or Bi-sulfosalts). The wall rock alteration is very pronounced with extensive brecciation and bleaching as well as pyrite and chalcopyrite impregnation. A chip sample of the vein exhibits raised values of gold (60 ppb) and bismuth (5 ppm). A chip sample of the hydrothermally altered zone also shows raised values of gold (30 ppb) and bismuth (30 ppm).

The area just west of Blengsdalen was also investigated on a reconnaissance basis. The area mainly consists of quartzitic and arkosic rocks, but an amphibolite dyke or sill has intruded the sediments. A minor area with widespread quartz-tourmaline-ankerite-pyrite veinlets was located (samples 2110073-74 - Sample Map 2). Analysis of these showed no gold enrichment. 200 m NE of this area abundant boulders of quartz-tourmaline veins (2-5 cm) were located (sample 2110075 - Sample Map 2). Analysis of these shows a content of 0.1 ppm Au, 7 ppm Bi, 3 ppm Ag and 620 ppm Cu.

Gjuv.

The location of the Gjuv Mine (occurrence) is shown in Fig.1. More precisely the old drift is located 100 m SE of the farm nedre Gjuv (sample 2110003 - Sample Map 10). According to Dons (1963), the occurrence was found before 1862. It has been described by Bagstevold (1918), Busvold (1919), Dokken (1943) and Pedersen (1984). In summary, the old workings (Fig.39) consist of an open pit with a lower 50 m long drift. The quartz-tourmaline-calcite veins have reported orientations of 0°-20°/ subvertical, 135°/weakly NE dip and 90°/30°-40°N which seems to reflect a complex vein pattern. The N-S veins seem to be the Thickest (> 1 m) whereas the NW-SE veins (up to four pa-



Fig.39. The mine dump and old workings of Gjuv.

parallel veins (\sim 30 cm) over a width of 15 m) seem to be thinner. The ore minerals are mainly bismuthinite and associated Bi-sulfosalts (cosalite, galenobismuthite etc.) and subordinate chalcopyrite, pyrite and gold. Pronounced hydrothermal wall rock alteration of the amphibolite similar to that of Bleka occurs.

The grade of the ore lenses seems to be 1-2% Bi, 5-10 ppm Au and minor Pb-Cu-Ag-Te. All of the mined vein ore has been thoroughly hand sorted and a total of 5-10 t of concentrated ore have been produced. This should amount to 300-500 kg of Bismuth.

The veins all occur in a 100 m thick amphibolite (the Kasin amphibolite) occurring parallel to and at a level above the Bleka amphibolite. It is possible that the complex vein pattern is a result of the position close to a major fault. The old workings are now more or less inaccessible, thus only a single grab sample from the mine dump was collected. It contains 0.16% Bi and 19 ppb Au.

3.5. Reconnaissance for Au-bearing veins in unknown areas.

Reconnaissance exploration for Bleka-type veins was also carried out in the Bleka amphibolite outside areas with known vein mineralization. The reconnaissance was mainly concentrated in areas with stream-sediment anomalies or in important lineament areas. A more or less continuous area of the Bleka amphibolite stretching from Skjesvatn in the NW to Hjartdal in the SE was traversed. In addition to the known vein occurrences, Bleka-type quartz-tourmaline veins were found in the area between Barstad and Bleka (Steinsruddalen - samples 2110083-85 - Sample Map 5), NE of Bleka (Ormemyre - sample 2110088 - Sample Map 6), in the area between Sverveli and Hjartdal (Mølledammen - sample 2110226 - Sample Map 2 and Vrestebakken - sample 2110225 - Sample Map 6), in the Hjartdal area and in the Blengsdalen area west of the known occurrences. Only the veins of Ormemyre and Steinsruddalen are of important size. At Steinsruddalen a 20 cm thick quartz-tourmaline vein ($55^{\circ}/70^{\circ}$ N) composed of quartz, very abundant tourmaline minor ankerite, pyrite and chalcopyrite is located. Very pronounced brecciation with quartz fragments in a tourmaline matrix as well as very strong hydrothermal wall rock alteration is observed.

Only the veins west of Blengsdalen are enriched in gold (0.1 ppm).

Apart from samples of Bleka-type veins, sampling of a major altered mylonite zone in the Bleka amphibolite at Brekka ytre in Ambjørndalen was performed. The mylonite zone which is 50-100 m wide and of unknown strike length exhibits pronounced epidotization, chloritization, calcite veining and pyrite impregnation. Bulk chip samples show low to raised gold contents (up to 29 ppb).

Reconnaissance for gold mineralization was also carried out in areas of acid volcanics and metasediments. No interesting alteration or mineralization was observed in the acid volcanics and no Bleka-type veins were seen in the metasediments. However, several quartz veins were chip sampled. Only one of the veins has gold values above the detection limit of 20 ppb.

This vein which has a gold content of 0.35 ppm is a 40 cm thick quartz vein with minor hematite. It occurs in quartzite and could be followed over more than 25 m (both ends open). It is located at Kåsastul north of Hjartdal (sample 2110017 - Sample Map 8).

4. DISCUSSION OF RESULTS.

The gold-bismuth veins of the Bleka Fold area occur geologically in a setting similar to known gold producing areas of the world. Some of the gold mines (e.g. the Sigma Mine) of the Val d'Or district, Abitibi Greenstone Belt of Canada are good examples as they are very similar to the Bleka-type veins.

The results of this investigation strongly indicate that all of the known occurrences of the Bleka Fold area (Bleka, Barstad, Nystaul, Blengsdalen and Gjuv) and the newly discovered Espelid vein swarm have a common ore genesis.

On the basis of the results from detailed investigations of the Bleka Mine area and the Espelid area a conceptual model for the Bleka-type veins of the Bleka Fold area is proposed (Fig.40).

During the Sweconorwegian orogeny, which was the last to affect the rocks of the Telemark area, the host rocks of the veins were tectonically deformed and metamorphosed to the amphibolite facies whereas the mineralized veins have not been subject to pervasive postore deformation and metamorphism. Thus, the veins were emplaced late in the geological history of the region.

As seen in Fig.40A the late folding around N-S to NNE-SSW axes in the area is a result of E-W to ESE-WNW max compression.

Assuming a level of the crust located somewhere close to the boundary zone between the ductile and the brittle regime (brittle-ductile regime) which is indicated by mapped tectonic structures (e.g. Fig.17) and by the metamorphic grade, a conjugate fracture system with an angle of 60°-90° would develop.

Actually, the lineament analysis shows a set of prominent ~ NE-SW and ~ NW-SE fractures to be present. The observed movement along the two prominent directions are also in agreement with the theoretical movements resulting from ESE-WNW compression. The NW-SE sinistral faults are the better developed

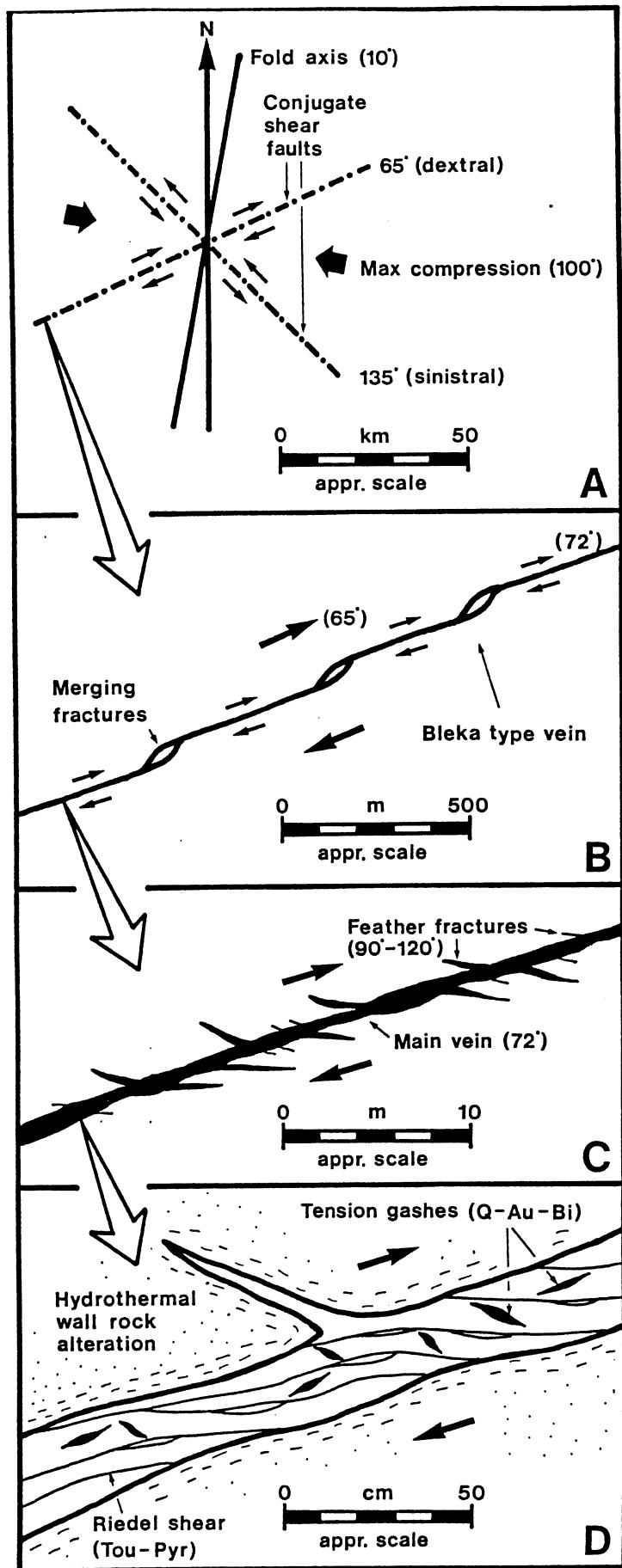


Fig. 40. Sketch of conceptual model for the Bleka-type veins of the Bleka Fold area.
A-D show various scales.

of the two conjugate fault directions in the central Telemark area and thus have taken up most of the shear movement. However, as seen in Fig.40B gold-bismuth bearing quartz-tourmaline-carbonate veins have been emplaced along the 65°-striking dextral shear zones (faults) where these intersect the mafic sills and dykes of the area. Bleka-type veins have not been located outside the amphibolites.

The individual dextral shear zones may be developed as shown in Fig.40B consisting of merging vein segments (fracture segments). From an economic point of view these areas of merging are very attractive because max tension and dilation occurs in these zones. A very strong support for this is seen in the enclosed Map 2 which shows the Bleka Main Vein System. It is evident that the major ore lenses occur in an area of merging between two en echelon veins. No significant gold-bismuth mineralization is observed outside this $300 \cdot 75 \text{ m}^2$ surface area. Assuming a depth of 150 m this implies that individual rock volumes potential for hosting high grade ore lenses are estimated to be in the order of 10 mill. t. Surface investigations indicate dilation in these zones in the order of 1-2%. In a more detailed scale (Fig.40C) the veins may exhibit pronounced pinch and swell structures and very characteristic feather fractures arranged parallel to max compression (100°) and perpendicular to max tension (Figs.35-37). These phenomena are a result of a slightly irregular fault plane and a shear movement in the brittle-ductile regime. As indicated, the shear movement is clearly dextral.

On a very detailed scale (Fig.40D) it can be observed how internal Riedel shear planes often expressed as tourmaline-sulphide veinlets occur in the main quartz vein (Fig.13). Theoretically, a set of late tension gashes should be developed. From the old reports of Kostke (1902) and Lauer (1943) it can be concluded that their observations on late cross-cutting quartz veins with visible gold and Bi-minerals can be related to these tension gashes. Other evidence of this is observed locally as open space filling of quartz crystals.

Pronounced hydrothermal alteration envelopes (Fig.40D) are developed throughout the area. Mineralogically, the amphibolite host rock increases its muscovite and calcite content towards the vein, whereas the hornblende and feldspar content decreases.

This results in an overall bleaching of the rock towards the vein and eventually, the altered amphibolite is transformed into a quartz-muscovite-calcite rock (Figs.5-8). This is indicative of percolating hydrothermal solutions causing the precipitation of the Au-Bi bearing quartz-tourmaline-carbonate veins in areas of max dilation. The mineral assembly of the veins: quartz-tourmaline-ankerite and subordinate calcite-dolomite-muscovite-epidote-chlorite-chalcopyrite-pyrite-bismuthinite-Bi-sulfosalts (cosalite?, galenobismuthite?, aikinite?, emblectite?)-native bismuth-native gold-maldonite?-bornite-galena-scheelite is indicative of a high temperature hydrothermal mineralization. In particular the symplectic intergrowths of gold and bismuth and the occurrence of maldonite indicate this. Zoning is observed both laterally and vertically. The lateral zoning is characterized by enrichment of tourmaline, carbonates and gold-bismuth minerals in the zones of merging (Fig.40B) and predominantly quartz with subordinate tourmaline and pyrite in veins outside these areas of max dilation. Vertical zoning is probably observed as introduction of galena at higher levels (Sverveli). The present relative levels of the known occurrences are so that the Espelid vein swarm occurs appr. 200 m above Bleka and Barstad and 300 m above Gjuv and Blengsdalen. As the veins were emplaced during the last period of major tectonic movements and assuming no later tilting there is reason to believe that the Espelid vein swarm contains Au-Bi mineralization at depth. This is further supported by the local enrichment of both elements which is interpreted as a halo phenomena.

The high temperature nature of the veins is probably due to underlying granite(s) which is exposed locally (Fjellstadfjell Granite N of Gjuv and a small granite plug SW of Bleka). Thus, the location of the Bleka-type veins is a result of regional tectonics combined with favorable host rock lithology (amphibolite) whereas the high temperature nature and to some degree the mineral assembly is controlled by underlying intrusions.

Impact on further exploration.

On a local scale the establishment of the conceptual model implies that zones of merging or other areas of max dilation are the main targets for gold-bismuth exploration. In the Bleka area one such zone ($300 \cdot 75 \text{ m}^2$ surface area) has already been outlined. The area just north of Sverveli could also represent the one end of such a zone, but this would have to be investigated in more detail. In the Espelid area which exhibits a large overall dilation the low exposure (5%) impedes detailed investigations. However, two areas (A - around samples 2110091-93, 95-98 and 100-115 and B - around samples 2110231-244 - Map 3) might represent zones of merging fracture segments.

Another impact on further exploration is the recognized Bi-halo. This means that future samples should be analyzed for both gold and bismuth.

On a larger scale it should also be noted that the Bleka-type veins seem to have a low, but widespread scheelite content. This seems to hold true for many gold-quartz veins in greenstones (e.g. Sigma Mine, Canada and Adelfors, Sweden). The occurrence of scheelite in particular in pan concentrates from areas of volcanic rocks should be checked for gold-quartz veins. This is as relevant for other areas in Norway as for the Telemark area.

Another point of interest is the result of the 1984 stream sediment sampling programme in central Telemark. In an area covering appr. 600 km^2 , 85% of the significant gold anomalies ($> 50 \text{ ppb}$) occur within a belt covering only 15% of the area. This belt which seems to be 5-10 km wide and has a strike similar to that of the Bleka veins runs from the Espelid-Bleka-Blengsdalen area in the SW to beyond Gjuv in the NE. The significance of this anomalous belt is uncertain, however, on a very large scale it could be postulated (mainly from interpretation of the 1:1,000,000 geological map) that a major fracture zone stretches all across the Southern Precambrian Province from the Stavanger area over Bykle-Dalen-Bleka-Kongsberg to the Oslo Rift and acted as a major channelway for hydrothermal solutions. However, this postulate is by no means supported by further investigations.

Economic potential.

The inferred reserves and metal productions from both the Bleka and the Gjuv Mines are small. The inferred original ore reserve of the Bleka Mine amounts to 50,000 t of vein material with a grade of 5 ppm gold or a diluted ore reserve of 250,000 t with 1 ppm gold. This includes a high grade part of ~10,000 t with 25 ppm, 15,000 t with 1 ppm and 25,000 t with appr. 0.1 ppm. Metal production includes 125 kg Au, 225 kg Ag, 20 t Bi and 70 t Cu from the early production period before year 1900 and 40 kg Au, 80 kg Ag, 6 t Bi and 10 t Cu from the 1935-39 production period.

Total metal production from ~7000 t of vein material:

Au - 165 kg

Ag - 300 kg

Bi - 25 t

Cu - 80 t

The figures for the Gjuv Mine are much more speculative, but only amount to 0.3-0.5 t of bismuth produced.

However, the actual economic potential is considered much better for the following reasons:

- a) the Bleka-type veins are known to contain ore lenses with very attractive grades - 25 ppm Au, 0.5% Bi, 45 ppm Ag and 1.5% Cu.
- b) known ore lenses occur in areas of merging between en echelon veins where max dilation occurred. The bulk volume of such zones is estimated to be in the order of 3-5 mill m³ = 10-15 mill t and no drilling has been performed.
- c) the located Espelid vein swarm represents a much larger hydrothermal system than the Bleka Main Vein System and exhibits a high overall dilation. 5-10 of the mapped veins exhibit average thicknesses of up to three times larger than the Bleka Vein. Assuming a similar grade the expected size of potential individual ore lenses of the Espelid vein swarm would be 50,000-100,000 t.

- d) bismuth and gold enrichment is observed locally within the Espelid vein swarm. This is believed to represent a halo phenomena from deeper-lying mineralization.
- e) assuming the observed high grade and minable widths for the ore lenses an ore reserve of only 150,000-200,000 t would be sufficient for a profitable small scale mining project (total revenue - 300-400 mill 1984-NoK).

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APPENDIX I.

ROCK SAMPLE DESCRIPTION LIST :

- 2110001 - Hydrothermally altered amphibolite with coarse-grained calcite/ ankerite veinlets with bismuthinite, chalcopyrite and pyrite. From Bleka main vein at Sverveli.
- 2110002 - Quartz vein with abundant chalcopyrite and minor pyrite and bismuthinite. Chlorite bearing. From upper Bleka mine dump.
- 2110003 - Quartz vein with pyrite and bismuthinite. Slickensides. From main mine dump at Gjuv.
- 2110004 - Quartz-ankerite vein with chalcopyrite, pyrite and bismuthinite. Chip sample from mine dump at Blengsdalen occurrence.
- 2110005 - Quartz vein with chalcopyrite and minor pyrite and bismuthinite. From mine dump at Blengsdalen occurrence.
- 2110006 - Quartz-ankerite vein with chalcopyrite, pyrite and bismuthinite. From mine dump at Blengsdalen occurrence.
- 2110007 - Quartz-ankerite-tourmaline vein with pyrite-chalcopyrite-bismuthinite. Rich ore from dump at Bleka.
- 2110008 - Quartz-ankerite vein with pyrite-chalcopyrite-bismuthinite. Rich ore from dump at Bleka.
- 2110009 - Quartz-carbonate vein with pyrite-chalcopyrite-bismuthinite? Rich ore from dump at Bleka.
- 2110010 - Quartz-tourmaline-ankerite vein with minor pyrite-chalcopyrite. Chip sample of mineralized vein material from main dump at Barstad occurrence.
- 2110011 - Quartz-tourmaline-ankerite vein with pyrite and minor chalcopyrite and bismuthinite? From main dump at Barstad occurrence.
- 2110012 - Quartz-tourmaline-ankerite vein with minor pyrite. Chip sample of 60 cm vein in the Barstad tunnel.
- 2110013 - Tourmaline-ankerite-quartz vein with abundant chalcopyrite. From mine level 610 at Bleka.
- 2110014 - Quartz-carbonate-tourmaline vein with chalcopyrite-bismuthinite veinlets. Rich outcropping ore from shaft level 595 at Bleka.
- 2110015 - Calcareous schist with pyrite pseudomorphoses. From Kivikdalen.
- 2110016 - Fine-grained amphibolite with quartz veinlets. From Øygården area.
- 2110017 - Quartz vein with minor hematite? From Kåsastul.
- 2110018 - Quartz-tourmaline vein with minor pyrite. From Espelid area.
- 2110019 - Hydrothermally altered amphibolite with pyrite and magnetite. Chip sample from Espelid area.

- 2110020 - Quartz vein with minor pyrite. Chip sample from Espelid area.
- 2110021 - Quartz-tourmaline-ankerite vein with minor pyrite and chalcopyrite. Chip sample from Espelid area.
- 2110022 - Hydrothermally altered amphibolite with pyrite and magnetite. Chip sample from Espelid area.
- 2110023 - Quartz-tourmaline vein with abundant pyrite and minor chalcopyrite. Sample from vein in Espelid area.
- 2110024 - Quartz-tourmaline vein with pyrite. Chip sample from Espelid area.
- 2110025 - Quartz-tourmaline-pyrite vein with tourmalinized wall rock (amphiolite). From Espelid area.
- 2110026 - Quartz-tourmaline-pyrite-wollastonite? veins. Chip sample of four minor veins in Espelid area.
- 2110027 - Hydrothermally altered wall-rock amphibolite with pyrite-magnetite tourmaline-apatite. Chip sample from Espelid area.
- 2110028 - Quartz-wollastonite vein. From Espelid area.
- 2110029 - Tourmaline-quartz vein with pyrite. From Espelid area.
- 2110030 - Quartz-tourmaline-ankerite vein with minor pyrite. Chip sample from Espelid area.
- 2110031 - Hydrothermally altered wall-rock amphibolite with pyrite. Chip sample from Espelid area.
- 2110032 - Silicified breccia with quartz and chlorite veinlets. Chip sample from Nystaul area.
- 2110033 - Quartz-tourmaline-ankerite vein with pyrite. Chip sample from Nystaul occurrence.
- 2110034 - Hydrothermally altered wall-rock amphibolite with pyrite. Chip sample from Nystaul occurrence.
- 2110035 - Quartz-tourmaline-ankerite vein with pyrite and euhedral magnetite octahedrons. Chip sample from Nystaul occurrence.
- 2110036 - Hydrothermally altered wall-rock amphibolite with pyrite. Chip sample from Nystaul occurrence.
- 2110037 - Quartz-hematite-sericite vein. From Nystaul area.
- 2110038 - Quartz-tourmaline-ankerite vein with minor pyrite. Chip sample from Nystaul occurrence.
- 2110039 - Coarse-grained hornblenditic amphibolite with disseminated pyrite. From Espelid-Nystaul area.
- 2110040 - Amphibolite with disseminated pyrite. From Espelid-Nystaul area.

- 2110041 - Amphibolite with disseminated pyrite. From Espelid-Nystaul area.
- 2110042 - Amphibolite with disseminated pyrite and magnetite. From Espelid area.
- 2110043 - Altered amphibolite with calcite-pyrite-chalcopyrite veinlet. From Haugstøl area.
- 2110044 - Quartz vein with abundant chalcopyrite. Chip sample from Haugstøl area.
- 2110045 - Tourmaline-quartz-ankerite vein with chalcopyrite and pyrite. Chip sample from Haugstøl area.
- 2110046 - Quartz vein with abundant chalcopyrite and minor bismuthinite. Chip sample from Bleka area.
- 2110047 - Chalcocite-calcite mylonite. From level 610 of Bleka mine.
- 2110048 - Chalcocite-calcite mylonite. From level 610 of Bleka mine.
- 2110049 - Chalcocite-calcite mylonite. From level 610 of Bleka mine.
- 2110050 - Chalcocite-calcite mylonite. From level 610 of Bleka mine.
- 2110051 - Chalcocite-calcite mylonite. From level 610 of Bleka mine.
- 2110052 - Chalcocite-calcite mylonite. From level 610 of Bleka mine.
- 2110053 - Ankerite-tourmaline vein with minor pyrite and chalcopyrite. From level 610 of Bleka mine.
- 2110054 - Quartz vein with minor tourmaline-ankerite-pyrite. Chip sample from level 550 of Bleka mine.
- 2110055 - Quartz vein with minor tourmaline-ankerite-pyrite-chalcopyrite. Chip sample from level 550 of Bleka mine.
- 2110056 - Quartz-tourmaline-ankerite vein with pyrite and chalcopyrite. Chip sample from level 525 of Bleka mine.
- 2110057 - Quartz-tourmaline-ankerite vein. From Bleka area.
- 2110058 - Quartz-tourmaline-ankerite vein with pyrite, chalcopyrite and bismuthinite. Chip sample from Blengsdalen area.
- 2110059 - Hydrothermally altered wall-rock amphibolite with abundant pyrite and chalcopyrite. Chip sample from Blengsdalen area.
- 2110060 - Quartz-tourmaline-ankerite vein with chalcopyrite and bismuthinite. From Bleka main vein.
- 2110061 - Ankerite-tourmaline vein with late quartz veinlets with abundant bismuthinite and chalcopyrite. One grain of native gold (1.1 mm) was observed in the quartz. From Bleka mine dump level 560.

- 2110062 - Quartz-tourmaline vein with abundant chalcopyrite and visible gold. Three to four grains of gold were observed (max 1 mm). From Bleka mine dump level 585.
- 2110063 - Quartz vein with minor tourmaline and pyrite. Chip sample from Bleka area.
- 2110064 - Quartz-tourmaline vein with minor pyrite. Chip sample from Bleka area.
- 2110065 - Quartz-tourmaline vein with minor pyrite. Chip sample from Bleka area.
- 2110066 - Quartz-tourmaline-ankerite vein with minor pyrite and chalcopyrite. Chip sample from Bleka area.
- 2110067 - Hydrothermally altered wall-rock amphibolite with pyrite and chalcopyrite. Chip sample from Bleka area.
- 2110068 - Quartz-pyrite-tourmaline vein. From Bleka area.
- 2110069 - Quartz vein with abundant chalcopyrite and minor pyrite and bismuthinite. From bleka area.
- 2110070 - Quartz-tourmaline vein. Chip sample from Bleka area.
- 2110071 - Quartz vein with minor feldspar. Chip sample from Bleka area.
- 2110072 - Quartz-tourmaline vein with minor pyrite. Chip sample from Bleka area.
- 2110073 - Tourmaline-quartz-carbonate-pyrite veinlets. Bulk chip sample from Blengsdalen area.
- 2110074 - Amphibolite with coarse-grained tourmaline veinlets. From Blengsdalen area.
- 2110075 - Quartz-tourmaline vein with pyrite and chalcopyrite. From Blengsdalen area.
- 2110076 - Quartz-pyrite vein in quartzitic host rocks. From Nutheim area.
- 2110077 - Quartz-ankerite vein with abundant chalcopyrite-pyrite and minor galena and bismuthinite. Chip sample from dump at Sverveli occurrence.
- 2110078 - Hydrothermally altered wall-rock amphibolite with pyrite and minor chalcopyrite. Chip sample from Sverveli area.
- 2110079 - Quartz-ankerite vein with pyrite-chalcopyrite and bismuthinite. From main vein at Sverveli.
- 2110080 - Quartz vein with abundant pyrite and galena. From Sverveli area.
- 2110081 - Quartz-ankerite vein with chalcopyrite, pyrite and galena. From Sverveli area.

- 2110082 - Quartz-tourmaline vein with chalcopyrite, galena and bismuthinite? Chip sample from Sverveli area.
- 2110083 - Quartz-tourmaline vein with minor ankerite and pyrite. Chip sample from Steinsruddalen area.
- 2110084 - Tourmaline-quartz-ankerite vein with minor pyrite and chalcopyrite. Chip sample from Steinsruddalen area.
- 2110085 - Hydrothermally altered wall-rock amphibolite with sericite. Extremely bleached. Chip sample from Steinsruddalen area.
- 2110086 - Quartz-tourmaline vein with minor pyrite. Chip sample from Livsengåsen area.
- 2110087 - Quartz-tourmaline vein with minor pyrite. Chip sample from Livsengåsen area.
- 2110088 - Brecciated quartz vein with minor tourmaline. Chip sample from Ormemyre area.
- 2110089 - Chlorite schist with ankerite and chalcopyrite-pyrite veinlets. From Blengsdalen area.
- 2110090 - Quartz vein with pyrite. In amphibolite sill. From north of Seljord along E76.
- 2110091 - Quartz-tourmaline-ankerite vein with abundant pyrite and minor chalcopyrite. Chip sample from Reiustaulhaugen area.
- 2110092 - Tourmaline-quartz vein with pyrite. Fine tourmaline-quartz banding. Chip sample from Reiustaulhaugen area.
- 2110093 - Quartz-tourmaline vein with abundant pyrite. Chip sample from Reiustaulhaugen area.
- 2110094 - Quartz-tourmaline vein with abundant pyrite. Chip sample from Reiustaulhaugen area.
- 2110095 - Quartz vein with minor tourmaline and pyrite. Chip sample from Reiustaulhaugen area.
- 2110096 - Quartz vein with minor tourmaline and pyrite. Chip sample from Reiustaulhaugen area.
- 2110097 - Quartz vein with minor tourmaline and pyrite. Chip sample from Reiustaulhaugen area.
- 2110098 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110099 - Quartz-tourmaline-pyrite vein. From Reiustaulhaugen area.
- 2110100 - Quartz vein with minor tourmaline and pyrite. Chip sample from Reiustaulhaugen area.
- 2110101 - Hydrothermally altered wall-rock amphibolite with abundant pyrite. Chip sample from Reiustaulhaugen area.

- 2110102 - Quartz vein with minor tourmaline and pyrite. Chip sample from Reiustaulhaugen area.
- 2110103 - Quartz vein with minor tourmaline and pyrite. Chip sample from Reiustaulhaugen area.
- 2110104 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110105 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110106 - Quartz vein with minor tourmaline and pyrite. Chip sample from Reiustaulhaugen area.
- 2110107 - Quartz vein with minor tourmaline and pyrite. Chip sample from Reiustaulhaugen area.
- 2110108 - Quartz vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110109 - Quartz-tourmaline-pyrite vein. Chip sample from Reiustaulhaugen area.
- 2110110 - Quartz vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110111 - Quartz vein with minor tourmaline and pyrite. Chip sample from Reiustaulhaugen area.
- 2110112 - Quartz vein with minor tourmaline and pyrite. Chip sample from Reiustaulhaugen area.
- 2110113 - Quartz vein with minor tourmaline and pyrite. Chip sample from Reiustaulhaugen area.
- 2110114 - Hydrothermally altered wall-rock amphibolite with abundant pyrite. Chip sample from Reiustaulhaugen area.
- 2110115 - Quartz vein with minor tourmaline and pyrite. Chip sample from Reiustaulhaugen area.
- 2110116 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110117 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110118 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110119 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110120 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.

- 2110121 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110122 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110123 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110124 - Hydrothermally altered wall-rock amphibolite with abundant pyrite. Chip sample from Reiustaulhaugen area.
- 2110125 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110126 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110127 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110128 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110129 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110130 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110131 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110132 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110133 - Quartz vein with minor tourmaline and pyrite. Chip sample from Reiustaulhaugen area.
- 2110134 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110135 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110136 - Quartz-tourmaline vein with minor pyrite. Chip sample from Reiustaulhaugen area.
- 2110137 - Quartz-tourmaline vein with minor pyrite. Chip sample from Bølskveven area.
- 2110138 - Hydrothermally altered wall-rock amphibolite with abundant pyrite. Chip sample from Bølskveven area.
- 2110139 - Quartz vein. Chip sample from Bølskveven area.

- 2110140 - Quartz vein. Chip sample from Bølskveven area.
- 2110141 - Quartz-tourmaline-pyrite vein with minor chalcopyrite.
Chip sample from Reiustaulhaugen area.
- 2110142 - Quartz-tourmaline-pyrite vein with minor chalcopyrite.
Chip sample from Reiustaulhaugen area.
- 2110143 - Quartz-tourmaline-pyrite vein. Chip sample from Reiustaulhaugen
area.
- 2110144 - Quartz vein. Chip sample from Nystaul area.
- 2110145 - Quartz-tourmaline vein with pyrite-chalcopyrite and bismuthinite?
From Reiustaulhaugen area.
- 2110146 - Mylonitized amphibolite with pronounced epidotization and chlo-
ritization. Pyritiferous. Chip sample from Brekka Ytre along E76.
- 2110147 - Mylonitized amphibolite with pronounced epidotization and chlo-
ritization. Pyritiferous and calcite bearing. Chip sample from
Brekka Ytre along E76.
- 2110148 - Mylonitized amphibolite with pronounced epidotization and chlo-
ritization. Pyritiferous and calcite bearing. Chip sample from
Brekka Ytre along E76.

- 2110201 - Quartz-tourmaline-ankerite vein with minor pyrite. From Nystaul area.
- 2110202 - Quartz-tourmaline-ankerite vein with minor pyrite. Chip sample from Nystaul area.
- 2110203 - Quartz-tourmaline-ankerite vein with minor pyrite. From Nystaul area.
- 2110204 - Quartz-tourmaline-ankerite vein with pyrite and chalcopyrite. Chip sample from Nystaul area.
- 2110205 - Quartz-tourmaline vein. Chip sample from Nystaul area.
- 2110206 - Quartz-tourmaline vein with pyrite. Chip sample from Nystaul area.
- 2110207 - Hydrothermally altered wall-rock amphibolite with pyrite and magnetite. Chip sample from Nystaul area.
- 2110208 - Quartz-tourmaline vein with pyrite. Chip sample from Nystaul area.
- 2110209 - Hydrothermally altered wall-rock amphibolite with pyrite and magnetite. Chip sample from Nystaul area.
- 2110210 - Quartz-tourmaline-ankerite vein with pyrite. From Reiustaulhaugen area.
- 2110211 - Quartz-tourmaline-ankerite vein with pyrite. Chip sample from Reiustaulhaugen area.
- 2110212 - Quartz-tourmaline vein with pyrite and mica. From Reiustaulhaugen area.
- 2110213 - Quartz vein. From Bølskveven area.
- 2110214 - Quartz vein with minor tourmaline-ankerite-muscovite and pyrite. From Bølskveven area.
- 2110215 - Quartz-tourmaline vein. Chip sample from Bølskveven area.
- 2110216 - Quartz vein with minor tourmaline. Chip sample from Bølskveven area.
- 2110217 - Hydrothermally altered wall-rock amphibolite with pyrite and magnetite. Extremely bleached. Chip sample from Bølskveven area.
- 2110218 - Quartz-tourmaline vein. Chip sample from Bølskveven area.
- 2110219 - Quartz vein with minor tourmaline. Chip sample from Bølskveven area.
- 2110220 - Quartz vein with slickensides. From Bølskveven area.
- 2110221 - Hydrothermally altered wall-rock amphibolite with slickensides. From Bølskveven area.
- 2110222 - Quartz-tourmaline vein with pyrite and chalcopyrite. Boulder from Livsengåsen area.
- 2110223 - Quartz veinlets. Chip sample from Langedal area.

- 2110224 - Quartz-sericite-chlorite-pyrile vein. Chip sample from Skvevsvatn area.
- 2110225 - Quartz-tourmaline vein. Chip sample from Vrestebakken area.
- 2110226 - Quartz-tourmaline vein with minor pyrite. Chip sample from Mølledammen area.
- 2110227 - Quartz vein with minor tourmaline. Chip sample from Bølskveven area.
- 2110228 - Hydrothermally altered wall-rock amphibolite with abundant pyrite. Chip sample from Bølskveven area.
- 2110229 - Quartz-tourmaline vein with minor pyrite. Chip sample from Bølskveven area.
- 2110230 - Hydrothermally altered wall-rock amphibolite with abundant pyrite. Chip sample from Bølskveven area.
- 2110231 - Quartz vein with minor ankerite. Chip sample from Bølskveven area.
- 2110232 - Quartz vein with minor tourmaline and sericite. Chip sample from Bølskveven area.
- 2110233 - Quartz-tourmaline vein. Chip sample from Bølskveven area.
- 2110234 - Hydrothermally altered wall-rock amphibolite with abundant pyrite. Chip sample from Bølskveven area.
- 2110235 - Hydrothermally altered wall-rock amphibolite with abundant pyrite. Chip sample from Bølskveven area.
- 2110236 - Quartz vein with minor sericite and ankerite. Chip sample from Bølskveven area.
- 2110237 - Hydrothermally altered wall-rock amphibolite with abundant pyrite. Chip sample from Bølskveven area.
- 2110238 - Quartz vein with minor tourmaline, sericite and ankerite. Chip sample from Bølskveven area.
- 2110239 - Quartz vein. Chip sample from Bølskveven area.
- 2110240 - Quartz-tourmaline-sericite vein. Chip sample from Bølskveven area.
- 2110241 - Quartz-tourmaline-pyrile vein. Chip sample from Bølskveven area.
- 2110242 - Quartz vein. Chip sample from Bølskveven area.
- 2110243 - Quartz vein with minor ankerite. Chip sample from Bølskveven area.
- 2110244 - Quartz vein with minor ankerite. Chip sample from Bølskveven area.
- 2110245 - Quartz vein with minor sericite. Chip sample from Bølskveven area.

2110246 - Quartz vein with minor sericite, chlorite and pyrite. Chip sample from Bølskveven area.

2110247 - Hydrothermally altered wall rock amphibolite with abundant pyrite. Chip sample from Bølskveven area.

2110248 - Quartz vein with minor tourmaline. Chip sample from Bølskveven area.

2110249 - Quartz vein with minor ankerite and sericite. Chip sample from Bølskveven area.

2110250 - Quartz-tourmaline-pyrite vein. Chip sample from Bølskveven area.

2110251 - Hydrothermally altered wall-rock amphibolite with abundant pyrite. Chip sample from Bølskveven area.

2110252 - Quartz vein with minor tourmaline and sericite. Chip sample from Bølskveven area.

2110253 - Quartz-tourmaline vein with minor pyrite. Chip sample from Bølskveven area.

2110254 - Quartz vein with tourmaline and pyrite. Chip sample from Bølskveven area.

2110255 - Hydrothermally altered wall-rock amphibolite with abundant pyrite. Chip sample from Bølskveven area.

APPENDIX II :

ANALYSIS RESULTS - LITHOSAMPLES.

SAMPLE NO.	Au PPB	Bi PPM	Ag PPM	Cu PPM	Pb PPM	Zn PPM	As PPM
2110001	73	6000	>250	240	-	-	-
2110002	-	-	-	-	-	-	-
2110003	19	1600	1	-	-	-	-
2110004	310	84	3	2760	-	-	-
2110005	103000	-	-	-	-	-	-
2110006	1800	-	-	-	-	-	-
2110007	5100	600	-	-	-	-	-
2110008	1350	2700	-	-	-	-	-
2110009	300	1000	-	-	-	-	-
2110010	<20	12	<1	412	-	-	-
2110011	30	12	-	-	-	-	-
2110012	<20	.3	<1	25	-	-	-
2110013	240	-	-	-	-	-	-
2110014	1600	2300	22	>10000	680	-	-
2110015	29	-	-	-	-	106	3
2110016	20	-	-	-	-	-	-
2110017	350	3	-	-	-	-	-
2110018	<10	-	-	-	-	-	-
2110019	<10	-	-	-	-	-	-
2110020	<10	-	-	-	-	-	-
2110021	<20	.2	<1	140	-	-	-
2110022	<20	-	-	-	-	-	-
2110023	15	<1	1	320	14	-	-
2110024	<10	-	-	-	-	-	-
2110025	-	-	-	-	-	-	-
2110026	<20	-	-	-	-	-	-
2110027	<20	-	-	-	-	-	-
2110028	-	-	-	-	-	-	-
2110029	-	-	-	-	-	-	-
2110030	<20	-	-	-	-	-	-

SAMPLE NO.	Au PPB	Bi PPM	Ag PPM	Cu PPM	Pb PPM
2110031	<10	-	-	-	-
2110032	<10	-	-	-	-
2110033	<20	.1	-	40	20
2110034	<20	-	-	-	-
2110035	<20	6.2	<1	60	30
2110036	<20	-	-	-	-
2110037	<10	-	-	-	-
2110038	37	-	-	-	-
2110039	-	-	-	-	-
2110040	-	-	-	-	-
2110041	-	-	-	-	-
2110042	-	-	-	-	-
2110043	<10	-	-	-	-
2110044	140	80	22	>10000	72
2110045	160	100	5	680	80
2110046	530	600	5	9200	236
2110047	<10	-	-	-	-
2110048	<10	-	-	-	-
2110049	<10	-	-	-	-
2110050	<20	-	-	-	-
2110051	60	-	-	-	-
2110052	<10	-	-	-	-
2110053	40	78	8	1050	-
2110054	<10	15	<1	23	-
2110055	<10	8	<1	40	-
2110056	2500	>1000	7	175	-
2110057	<20	-	-	-	-
2110058	60	5	<1	62	12
2110059	30	30	-	81	-
2110060	1680	72	2	135	66
2110061	68000	>1000	>250	6900	>10000
2110062	94000	>1000	41	5300	>10000
2110063	<20	-	-	-	-
2110064	<20	-	-	-	-
2110065	<20	-	-	-	-

SAMPLE NO.	Au PPB	Bi PPM	Ag PPM	Cu PPM	Pb PPM
2110066	<20	52	3	800	66
2110067	<20	-	-	-	-
2110068	70	260	19	1930	270
2110069	770	244	43	>10000	282
2110070	<20	-	-	-	-
2110071	<20	-	-	-	-
2110072	<20	-	-	-	-
2110073	<10	-	-	-	-
2110074	-	-	-	-	-
2110075	100	7	2	620	36
2110076	<20	-	-	-	-
2110077	115	126	73	3790	910
2110078	<20	-	-	-	-
2110079	280	>1000	89	5000	>10000
2110080	80	244	162	118	>10000
2110081	<10	22	12	2660	502
2110082	300	292	221	5200	>10000
2110083	<20	-	-	-	-
2110084	<20	-	-	-	-
2110085	<20	-	-	-	-
2110086	<20	-	-	-	-
2110087	<20	-	-	-	-
2110088	<20	-	-	-	-
2110089	<20	20	3	2300	36
2110090	87	-	-	-	-
2110091	29	.4	1	78	12
2110092	<20	-	-	-	-
2110093	<20	-	-	-	-
2110094	<20	-	-	-	-
2110095	<20	-	-	-	-
2110096	<20	-	-	-	-
2110097	<20	-	-	-	-
2110098	<20	-	-	-	-
2110099	38	-	-	-	-
2110100	<20	.7	1	8	8

SAMPLE NO.	Au PPB	Bi PPM	Ag PPM	Cu PPM	Pb PPM
2110101	<20	-	-	-	-
2110102	<20	-	-	-	-
2110103	<20	-	-	-	-
2110104	<20	-	-	-	-
2110105	<20	-	-	-	-
2110106	<20	-	-	-	-
2110107	<20	-	-	-	-
2110108	<20	-	-	-	-
2110109	<20	-	-	-	-
2110110	<20	-	-	-	-
2110111	<20	-	-	-	-
2110112	<20	-	-	-	-
2110113	<20	2	<1	40	8
2110114	<20	-	-	-	-
2110115	<20	-	-	-	-
2110116	<20	-	-	-	-
2110117	<20	-	-	-	-
2110118	<20	-	-	-	-
2110119	<20	-	-	-	-
2110120	<20	-	-	-	-
2110121	<20	1	<1	9	6
2110122	<20	-	-	-	-
2110123	<20	-	-	-	-
2110124	<20	-	-	-	-
2110125	<20	-	-	-	-
2110126	<20	-	-	-	-
2110127	<20	-	-	-	-
2110128	<20	-	-	-	-
2110129	<20	-	-	-	-
2110130	<20	-	-	-	-
2110131	<20	-	-	-	-
2110132	<20	-	-	-	-
2110133	<20	-	-	-	-
2110134	<20	-	-	-	-
2110135	<20	-	-	-	-

SAMPLE NO.	Au PPB	Bi PPM	Ag PPM	Cu PPM	Pb PPM
2110136	<20	.7	<1	58	4
2110137	<20	.9	<1	18	4
2110138	<20	-	-	-	-
2110139	<20	-	-	-	-
2110140	<20	-	-	-	-
2110141	28	24	3	242	18
2110142	28	55	1	32	56
2110143	<20	-	-	-	-
2110144	84	-	-	-	-
2110145	<20	144	3	10	140
2110146	<20	-	-	-	-
2110147	<20	-	-	-	-
2110148	29	-	-	-	-
2110201	<10	-	-	-	-
2110202	90	-	-	-	-
2110203	210	-	-	-	-
2110204	80	80	2	210	46
2110205	55	-	-	-	-
2110206	<10	-	-	-	-
2110207	<10	-	-	-	-
2110208	<10	-	-	-	-
2110209	50	-	-	-	-
2110210	<10	-	-	-	-
2110211	<10	2	<1	35	16
2110212	<10	2	<1	5	6
2110213	15	-	-	-	-
2110214	15	1	<1	5	6
2110215	<10	-	-	-	-
2110216	<10	-	-	-	-
2110217	<10	-	-	-	-
2110218	<10	-	-	-	-
2110219	<10	-	-	-	-
2110220	<10	-	-	-	-

SAMPLE NO.	Au PPB	Bi PPM	Ag PPM	Cu PPM	Pb PPM
2110221	-	-	-	-	-
2110222	<10	4	2	490	46
2110223	20	-	<1	-	-
2110224	<20	-	-	-	-
2110225	<32	2	<1	45	20
2110226	<20	-	-	-	-
2110227	<20	.2	<1	25	6
2110228	<20	-	-	-	-
2110229	<20	-	-	-	-
2110230	<20	-	-	-	-
2110231	<20	-	-	-	-
2110232	<20	<.1	<1	29	6
2110233	<20	-	-	-	-
2110234	<20	-	-	-	-
2110235	<20	-	-	-	-
2110236	<20	-	-	-	-
2110237	<20	-	-	-	-
2110238	<20	-	-	-	-
2110239	<20	<.1	<1	8	4
2110240	<20	-	-	-	-
2110241	28	-	-	-	-
2110242	<20	-	-	-	-
2110243	<20	-	-	-	-
2110244	<20	-	-	-	-
2110245	<20	-	-	-	-
2110246	<20	<.1	<1	8	4
2110247	<20	-	-	-	-
2110248	<20	-	-	-	-
2110249	<20	-	-	-	-
2110250	<20	-	-	-	-
2110251	<20	-	-	-	-
2110252	<20	.4	<1	72	4
2110253	<20	-	-	-	-
2110254	<20	-	-	-	-
2110255	<20	-	-	-	-

LITHOGEOLOGI HOVIN KOBBERVERK, TELLEVÆRK.

Indledning.

På basis af feltarbejde og påfølgende mikroskopering (Schønwandt, 1983 & Schønwandt, 1984) blev det besluttet at foretage en mindre feltaktion i 1984 med henblik på at etablere et lithogeokemisk kort over amfiboliten (malmzone-amfiboliten i tidligere rapporter) ved Hovin Kobberverk.

Idéen med at foretage den lithogeokemiske kortlægning var primært:

- i) At forsøge at verificere det på gamle analyser etablerede Cu-anomalimønster i amfiboliten (kort 4 i Schønwandt, 1983).
- ii) at undersøge om amfiboliten ved Hovin Kobberverk var abnormal i metal-indholdet i forhold til de omkringliggende amfiboliter.

I anden række at få bragt klarhed over feldspat-porfyrrens relation til amfiboliten, idet mikroskoperingen havde sandsynliggjort, at denne bjergart ikke havde det i feltrapporten (Schønwandt, 1983) foreslæde slægtskab til amfiboliten.

Geologisk kortlægning.

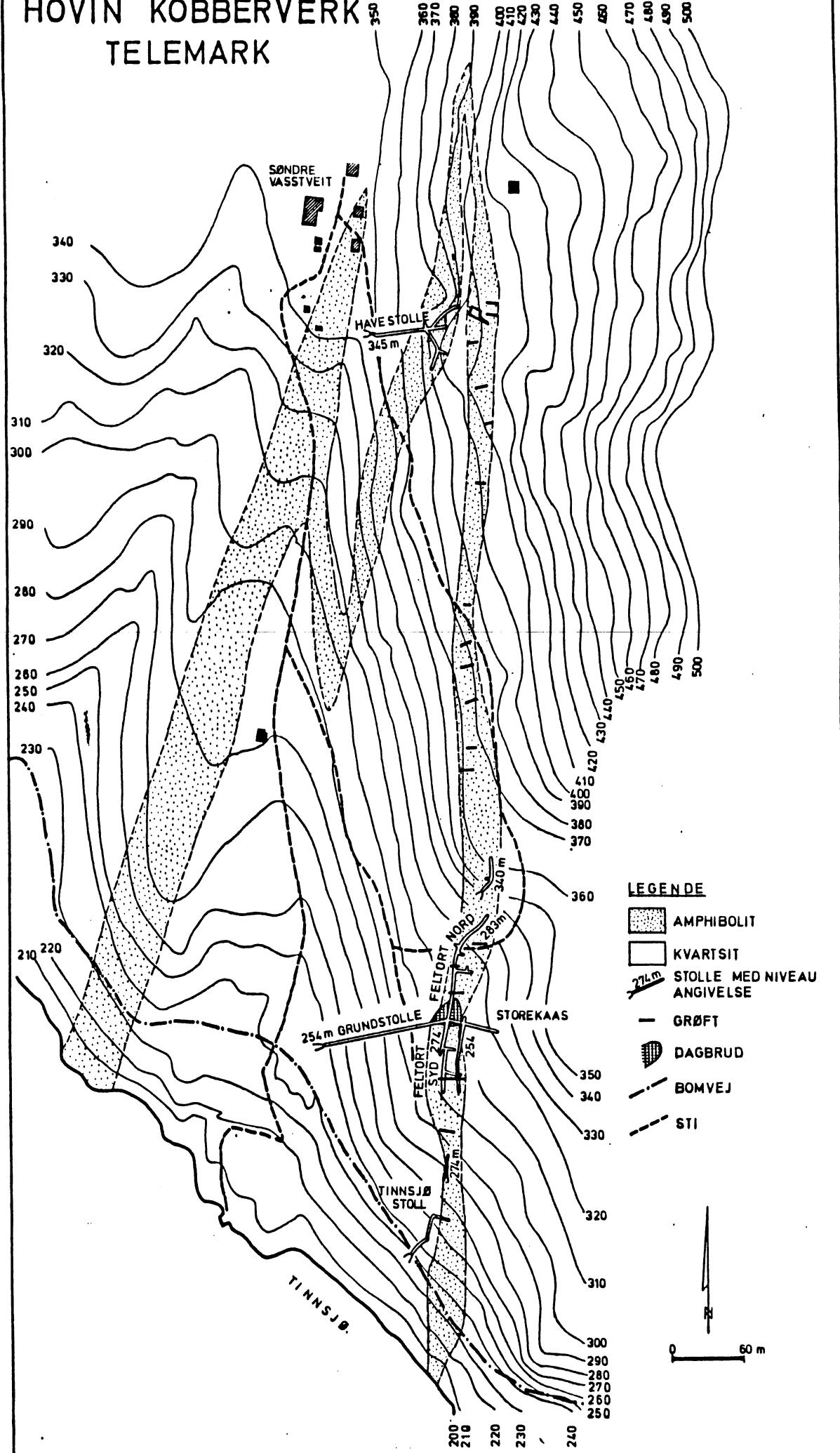
Den detaljerede kortlægning har vist, at de tidlige amfibolitområder omkring Hovin Kobberverk (malmzone-amfiboliten (østligst), Havestolle-amfiboliten og imprægnations-amfiboliten (vestligst)) udgør én og samme amfibolit (fig. 1). Amfiboliten, der ligger i en kvartsit sekvens, er isoklinalt foldet og har form som en zik-zak fold.

Alle målte småfolder har et N-S-gående eksialplan, der hælder stejlt mod øst (ca. 85°) og en mod nord dykkende (ca. 30°) foldeakse. På basis af småfolderne mener man formode, at zik-zak folden består af en østlig og en vestlig antiform med en mellemliggende synform.

Amfibolitten varierer i tynde fra 10-15 m op til ca. 50 m. Ombygningen i den østlige og vestlige anti-

HOVIN KOBBERVERK

TELEMARK



form ligger henholdsvis 1100 m og 800 m i strøgretning fra Finnsjøen (fig. 1).

Lithologisk er der en variation i amfiboliten. Den østlige flanke af zik-zak folden består helt overvejende af en amfibol-porfyrblastisk amfibolit med varierende konkordante zoner af biotit-førende amfibolit. Der er både skarp og glicende kontakt mellem disse to bjergarter.

I den nordlige trediedel af den østlige flanke glider amfiboliten jævnt over i en melanokratisk biotit-gneis. Biotit-genisen har konkordante zoner af varierende bredde af amfibolit og/eller porfyrblastisk amfibolit.

Den resterende del af zik-zak folden domineres af den melanokratiske biotit-gneis, der dog stedvis bliver mindre melanokratisk. I den sydligste del af zik-zak foldens vestflanke bliver amfiboliten igen den dominerende bjergart.

Den eneste bjergart inden for amfiboliten, der klart skiller sig ud er dioriten i dagbruddet Storekaas' østlige del. Dioriten fører, som tidligere omtalt (Schönwandt, 1983 & 1984), dissemineret chalcocit og bornit. Til trods for at dioriten fører disseminerede sulfider svarende til ca. 1% Cu og ca. 100 ppm Ag, så er denne bjergart ikke blevet betragtet som malm, da brydning fundt sted i århundredets begyndelse, idet bjergart i det store og hele er blevet ladt tilbage i dagbruddet.

Dioriten skennes at have udgjort et ellipseformet areal med en storakse på ca. 20 m og en lilleakse på ca. 7-8 m. Storaksen i ellipsen stryger næsten N-S og er således konform med amfiboliten. Dioriten er i kontaktzonen til amfiboliten kraftig smaret og fører et gneisagtigt dissemine. Dette medfører, at der er en glicende overgang mellem den gneisagtige diorit og den smerede amfibolit. Den smerede kontakt samt dioritetsmønstre vedje har begrænset yderligere

information om relationen mellem amfiboliten og dioriten.

En fornyet undersøgelse af de to feldspat-porfyr lokaliteter i det stærkt overdækkede område nord for dagbruddet har sandsynliggjort, at det drejer sig om løsblokke.

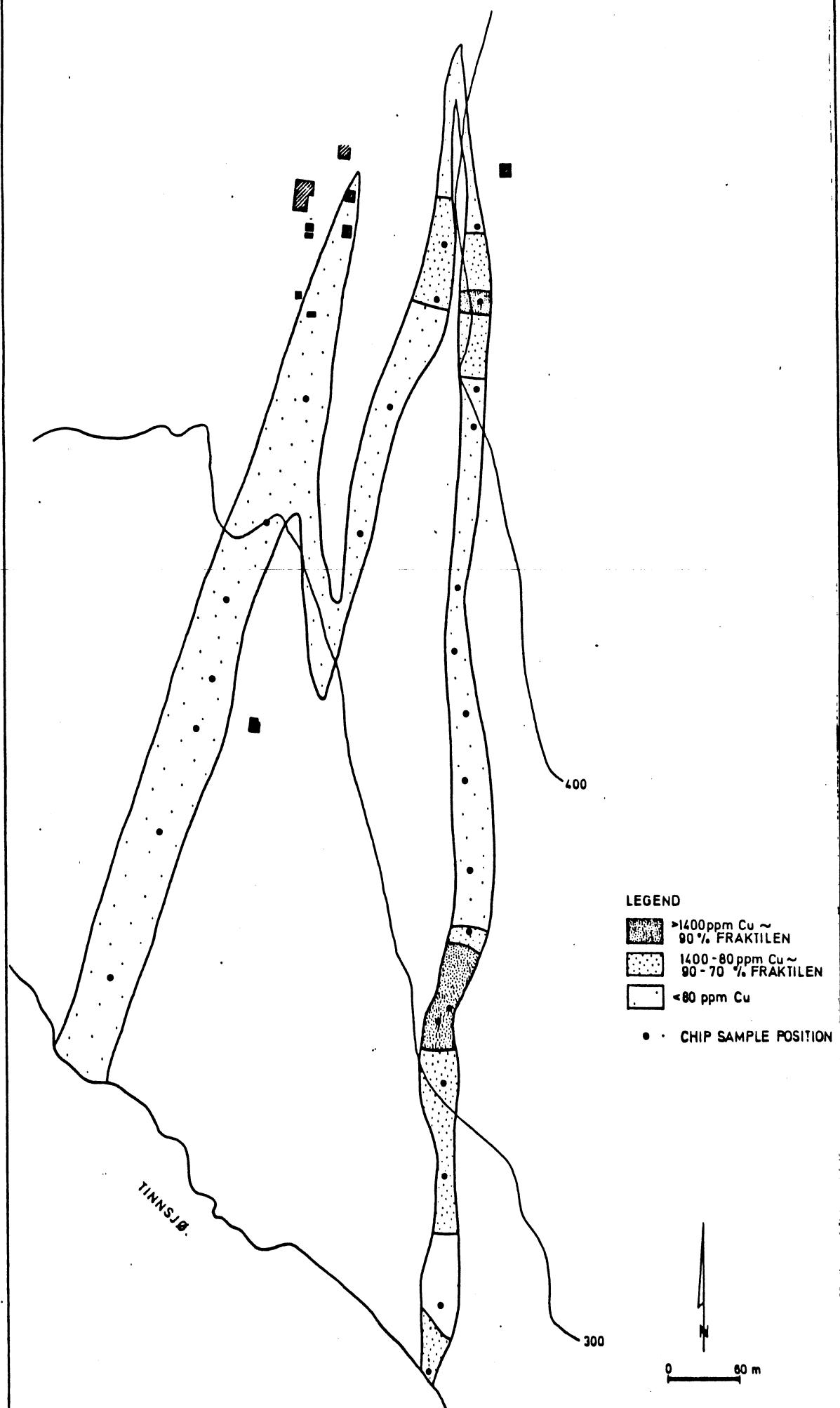
Sammenfattende kan man herefter om geologien ved Hovin Kobberverk sige, at mineraliseringen er knyttet til en amfibolit, der ligger i en kvartsit sekvens. Amfiboliten er foldet og har et vertikalt N-S-gående aksialplan. Baseret på småfolder antages det, at foldeaksen dykker ca. 30° mod nord. Folden har zik-zak form og består af en østlig og vestlig antiform med en mellemliggende synform. Kinematisk repræsenterer deformationen en shear-folding. Den lithologiske variation inden for amfiboliten, repræsenteret ved bjergarter fra porfyroblastisk amfibolit til melanokratisk biotitungneis, repræsenterer formodentlig mere et resultat af deformation og metamorfose end en oprindelig lithologisk variation i amfiboliten.

Dioriten i Storekaas dagbrud skiller sig ud fra de andre bjergarter i amfibolitkomplekset ved bl. a. strukturelt ikke at være konform med amfibolitten i mesoskala, hvorimod den i handstyksskala er konform med amfibolitten på grund af kontakt-zonens sterke shear deformation.

Lithogeokemi.

Der blev indsamlet 27 chip samples rimeligt fordelt over den zik-zak foldede amfibolit. Prævepunkterne er markeret med en prik på fig. 2, men repræsenterer i den udstrekning blotnægtsgraden tilledt, et tværprofil af amfibolitten. Alle chip samples er epigenetiske mineraliseringer i form af kvarts-kalcit og unagtet prævesterrrelsen varierer alt uafhængig af amfibolittens brude- og/eller blotningsgrad fra ca. 1 til 5 kg. prævestoren selten har været prospektøringsafdelings største procedure for lithoprøver og prævne

LITHOGEOKEMISK KORT OVER AMFIBOLITEN VED HOVIN KOBBERVERK TELEMARK



er blevet analyseret for Cu og Ag (tabel 1).

Fig. 3 viser Cu-analyserne plottet på sandsynligheds-papir. Den beskedne prøvemengde træt i betragtning, så kan plottet tolkes som repræsentende to lognormale fordelinger. Der er her valgt at lade de to bøjninger i kurveforlebet ved 80 og 1400 ppm Cu repræsentere henholdsvis tærskelverdiens for amfibolitens normale Cu-indhold og den nedre grænse for hvad der kan betragtes som diorit-mineraliseringen og dens umiddelbare mineraliserings-aureole. Værdierne mellem 80 og 1400 ppm repræsenterer i øvrigt forstånd den mineraliserings-aureole, der ligger omkring dioriten.

Værdierne på 80 og 1400 ppm svarer til henholdsvis 70% og 90% fraktilen. Den fundne baggrundsværdi for Cu i amfiboliten svarer til niveauet i tilsvarende bjergarter andre steder.

Chip sample analysen af dioriten (større end 1,6 Cu og 86 ppm Ag) svarer til hvad man kunne forvente på basis af tidligere analyser af godt mineraliserede enkelt prøver. Vigtigt i denne sammenhæng er også at analyseresultatet bekræfter den tidligere påpegede korrelation mellem Cu- og Ag-indholdet i chalcocit-bornit paragenesen (Schenwendt, 1985).

TABEL 1

Prøve nr.	ppm Cu	ppm Ag	Prøve nr.	ppm Cu	ppm Ag
2110370	360	1	2110371	27	1
2110372	10.000	86	2110373	2080	8
2110374	480	1	2110375	1.25	1
2110376	80	1	2110377	.45	1
2110378	.1	1	2110379	30	1
2110380	55	1	2110381	55	1
2110382	1.150	1	2110383	5	1
2110384	15	1	2110385	550	4
2110386	.9	1	2110387	1.510	9
2110388	.1	1	2110389	50	1
2110390	52	1	2110391	47	1
2110392	44	1	2110393	46	1
2110394	52	1	2110395	15	1
2110396	55	1			

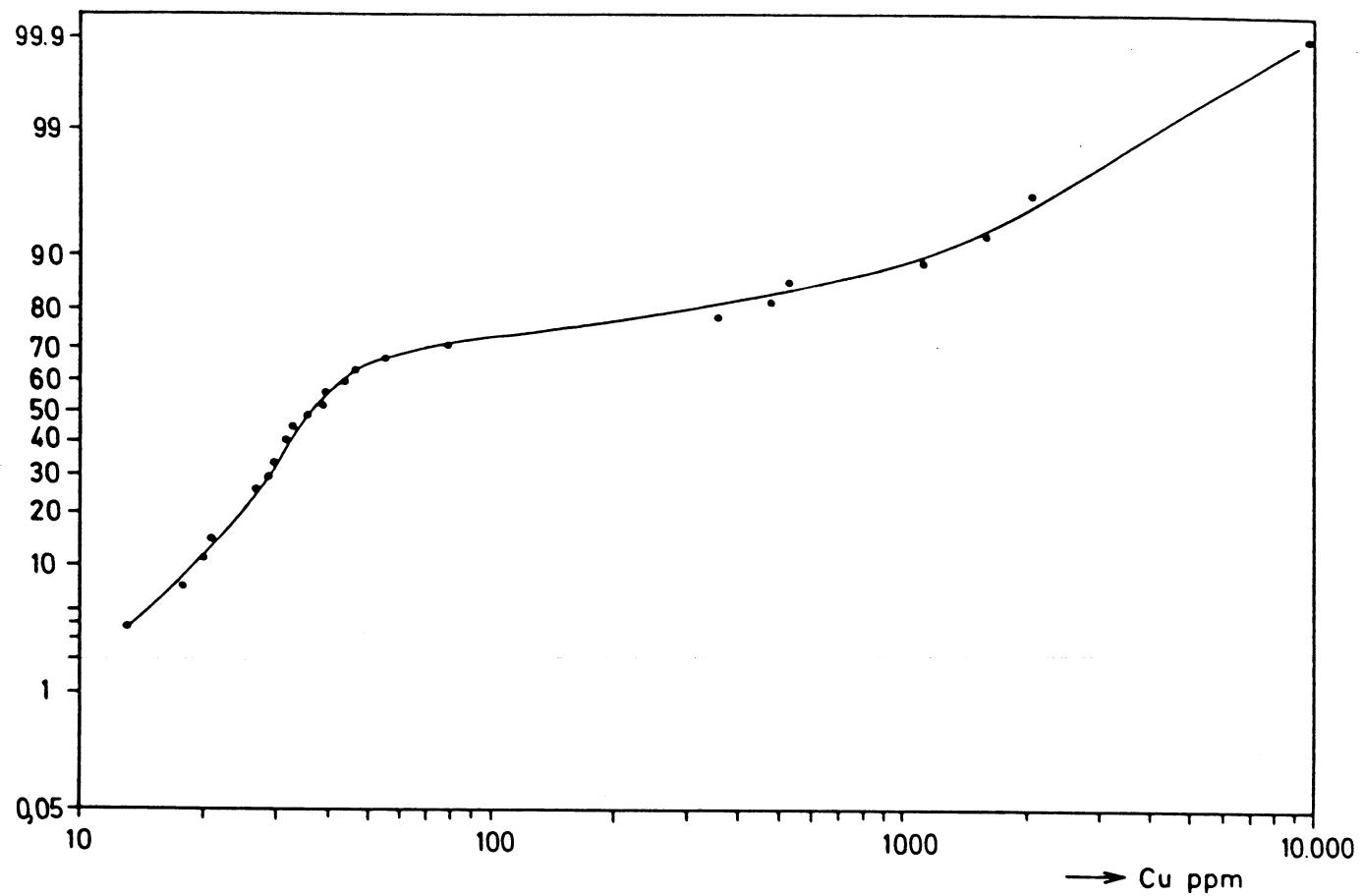


FIG. 3. FOR NÆRMERE FORKLARING SE TEKST.

Analyseresultaterne er plottet på kort (fig. 2). Kortet viser fordelingen af prøverne over 1400 ppm Cu (svarende til mere end 90% fraktilen) og mellem 1400 og 80 ppm Cu (svarende til 90%-70% fraktilen) samt prøver med et Cu-indhold under 80 ppm (svarende til amfibolitens baggrundsverdi). Grænsedragningen mellem de forskellige Cu-intervaller er fundet ved linjeforudsigelse.

Af det lithologiske kort fremgår, at der inden for zik-zak foldens østlige flanke findes to anomale områder i) et mod syd omkring Storekæs dagbrud og ii) et mod nord omkring Have stollen. Begge anomaliområder er opbygget op af flere anomalie prøver og i denne tilfælde er anomaliområdet zonart opbygget med de højeste verdier i den centrale del af anomlien. For det nordlige anomaliområdet vedkommende strækker den ydre aureole sig også over på den østlige antiforms vestlige flanke.

Det lithogeokemiske kort (fig. 2) verificerede de to anomaliområder omkring Storekæs og Have stollen, som var blevet etableret på basis af gennem analyser. Derimod kunne det tredje anomaliområde (se fig. 4 i Schenck, 1963) ikke bekræftes. Det lithogeokemiske kort viser også, at amfibolit i zik-zak foldens østlige flanke (tidligere rapportens malmonzon-amfibolit) ikke er normal i Cu-indholdet i forhold til amfiboliten i den øvrige del af zik-zak folden. Indviduelle anomaliområderne optimerer som værligefinere de partier inden for en i Cu-indhold lavrigt normal amfibolit. Dette lithogeokemiske billede understyrer yderligere antagelsen af dioritten som dræng til det konstituerende mineraliseringssystemet (benominert, 10%).

Det plante-geologiske udvalg består af 10 prøver fra zik-zak foldens østlige flanke, 10 prøver fra zik-zak foldens vestlige flanke og 10 prøver fra den østlige amfibolitzone i området omkring Storekæs (fig. 2).

Analyseresultaterne er samlet i tabellen, tabel 2.

at chalcocit-bornit findes dissemineret i dioriten og at denne mineralisering er omgivet af en epigenetisk mineraliseringsaureole hvis paragenese endre sig med afstanden fra dioriten. Nærmest dioriten består paragenesen af bornit-chalcocit, dernest kommer en zone med bornit-kobberkis og yderst en ren kobberkis paragenese. Mineraliseringerne er næsten hundrede procent knyttet til amfibolit komplekset.

Ser man på anomalimønsteret omkring Storekaas (fig. 2) så viser det, at anomalien er trukket ud i sydlig retning. Denne skævhed kan have mange orsager, som f.eks.

- i) At den epigenetiske mineraliseringsaureole ikke er koncentrisk omkring primærmineraliseringen (dioriten)
- ii) at en begrænset præcisenhed kombineret med at linierinterpolation har været anvendt ved grænsedragningen mellem de forskellige Cu-intervaller
- iii) en topografieffekt d.v.s. primærmineraliseringens forløb i forhold til terrænet.

Man kan forestille sig andre orsager, men de nævnte eller en kombination af disse skænkes at have været en dominerende årsag til skævheden i anomalimønsteret.

Først var iore disse faktorer er der fremstillet et N-S-gående profil i amfiboliten fra Tinnsjøen til lit nord for Storekaas (fig. 4). I dette profil er vist udstrekningen af de lithogeokemiske aureoler samt Cu-profilet igennem aureolerne. Endvidere er vist støller og tverslag fra 54 m, 74 m samt 93 m niveauerne.

Sammenholder man de gængede Cu-analyser fra støller og tverslag (Schmwandt, 1955 fig. 3) med de fundne lithogeokemiske aureoler, so forestrives indtrykket af at aureolene ikke er dygtige til at udvikle en Cu-verdelomme i den ydede præcision, hvilket i øvrigt er stor, at de godt kunne repræsentere den mineraliserede diorit.

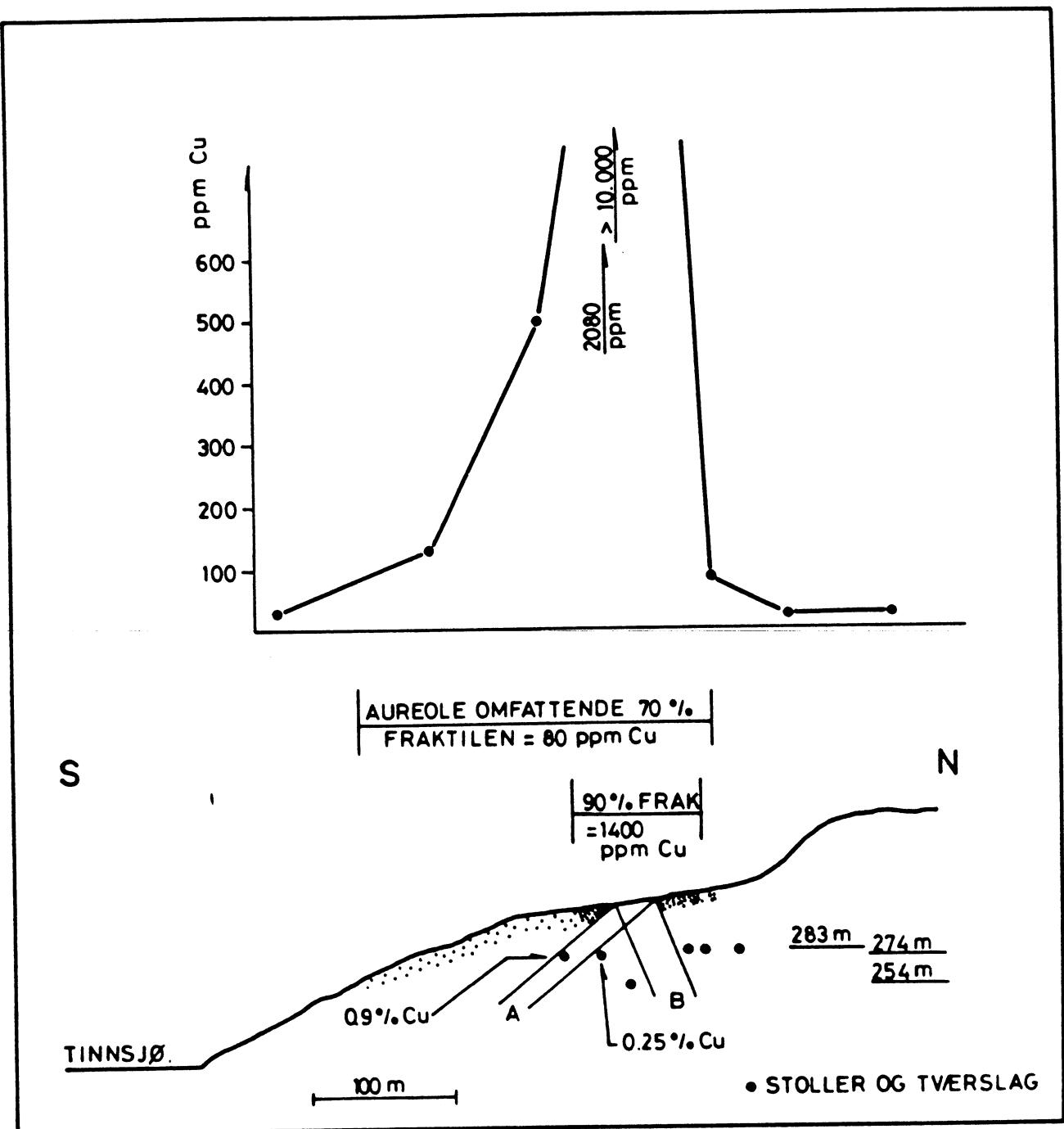


FIG. 4 , FOR NÆRMERE FORKLARING SE TEKST.

Det skal her inskydes, at de gamle analyser fra stoller og tværslag (fig. 3 i Schenckhardt, 1985) viser højere værdier, end de der er fundet ved denne undersøgelse. Dette henger formodentlig sammen med, at epigenetiske årer bevidst er undgået i denne undersøgelse. Man kan derfor ikke bruge de gamle analyser til at komplimentere denne undersøgelses aureolemønster med. Imidlertid må så store værdier, som det sydligste tværslag i niveau 274 m viser, enten henvores til dioriten eller til en usædvanlig stor koncentration af epigenetiske årer.

Diskussion.

Med de oplysninger der på nuværende tidspunkt er tilgengeligt så er det muligt, at dioriten har det mod syd dykkende forleb ~~som er indikeret med A~~ på fig. 4. Et sådant forleb af dioriten vil i meget høj grad give en tilfredsstillende forklaring på det fundne aureolemønster, når man tager topografiens og prævpunkterne i betragtning.

Det kan imidlertid ikke udelukkes at dioriten har det med B indikerede forleb. Et sådant forleb, med stejlt dyk mod nord, vil imidlertid ikke umiddelbart kunne forklare aureolemønsteret uden samtidig at acceptere en relativ stor prisær ikkeheds i dette mønster.

Det samme problem vil man ryge ind i, hvis man antager at dioriten er et meget lokalt fenomen, der ikke strækker sig mange 10-tals meter under overfladen.

Ett relativt fladt dyk af dioriten mod nord, vil også kunne forklare aureolemønsteret, hvis man samtidig antager, at dioriten i storkest udgør en lille rest af et større dioritfelt, der i det store og hele er ortoklastisk, fast i lysgrått i niveau 133 ikke tillader dioriten at formere sig i nogen form.

Alle områder med 100 m højdevariancer omstillerne for- alderstypen, så dioriten i fig. 4 på kontakten mellem diorit og ortoklastisk diorit følger en relativt

flanke. Det er intet der tyder på, at dette ikke skulle være tilfældet. Denne specielle position kan umiddelbart forklæres på idet mindste to måder:

- i. Dioriten, der som tidligere omtalt petrografisk er nært beslagtet med amfiboliten, kan være et differentiat af amfiboliten efter dens enplacering. Dette vil, amfibolitens størrelse taget i betragtning, resulterer i en meget beskeden størrelse af dioriten.
- ii. Dioriten repræsenterer en selvstændig intrusion, men fra det samme magmakammer. Denne mulighed vil kunne give en betragtelig størrelse på dioriten.

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Størrelsen på den lithogeokemiske aureole tyder på, at dioriten har/eller har haft en betragtelig størrelse, hvorfor det er mest sandsynlig at dioriten repræsenterer en selvstændig intrusion.

Forløbet af denne intrusion set i lyset af aureole-monsteret, kan enten være et relativt fladt dyk mod syd eller nord. Er dykket forløbet mod nord, så må hovedparten af dioriten være bort-eroderet. Hvis dykket varer mod syd, så har dioriten formodentlig et forløb nær det, der er indikeret med A på fig. 4.

Konklusion og udeafslutninger.

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Den lithogeokemiske undersøgelse tyder på, at dioriten repræsenterer en selvstændig intrusion, der er nært beslagtet med amfiboliten.

Det er næst sandsynlig at dioriten har et dyk mod syd nogenlunde svarende til det, der er indikeret med A på fig. 4. Et fladt dyk mod nord er mindre sandsynligt på grund af de høje Cu-værdier i de sydlige tværsag i niveau 274 m.

En mere detaljeret lithopetrografisk undersøgelse på overfladen vil sandsynligvis have bringt størrelsen på den øvre dioritem forløb, ved hvilken epigenetiske omformingsaureole en forvandlet havde en del "grønligk stål" og også formet blottningerne er ringe.

Det kan derimod anbefales :

at foretage en kortlægning og prøveindsamling af niveau 274 m. Adgang til niveauet vil formodentlig kunne ske gennem skakten i Storekaas dagbrud. Niveau 274 m ligger ca. 34 m under dagbrudsniveauet.

at det overvejes hvilke geofysiske metoder, der bedst vil kunne afsløre om dioritten har det med A indikerede forløb i fig. 4.

at økonomiske overslag foretages på basis af en malmværdi på 1,0 Cu og 90 ppm Ag med henblik på at få en minimums størrelse på malmten. Dette tal bør indgå i de geofysiske overvejelser.

Når prospekteringsafdelingens medarbejdere har foretaget de geofysiske og økonomiske overvejelser vil det være nønsigtsmæssigt, at der i fællesskab udarbejdes et notat vedrørende disse overvejelser.

Referencer.

Schönwandt, 1983 : Feltrapport over Hovin Kobberverk, Telemark. Intern Hydro Rapp.

Schönwandt, 1984 : Notat vedrørende mikroskopering af tyndslib og polerprover fra Hovin Kobberverk, Telemark. Intern Hydro Rapp.

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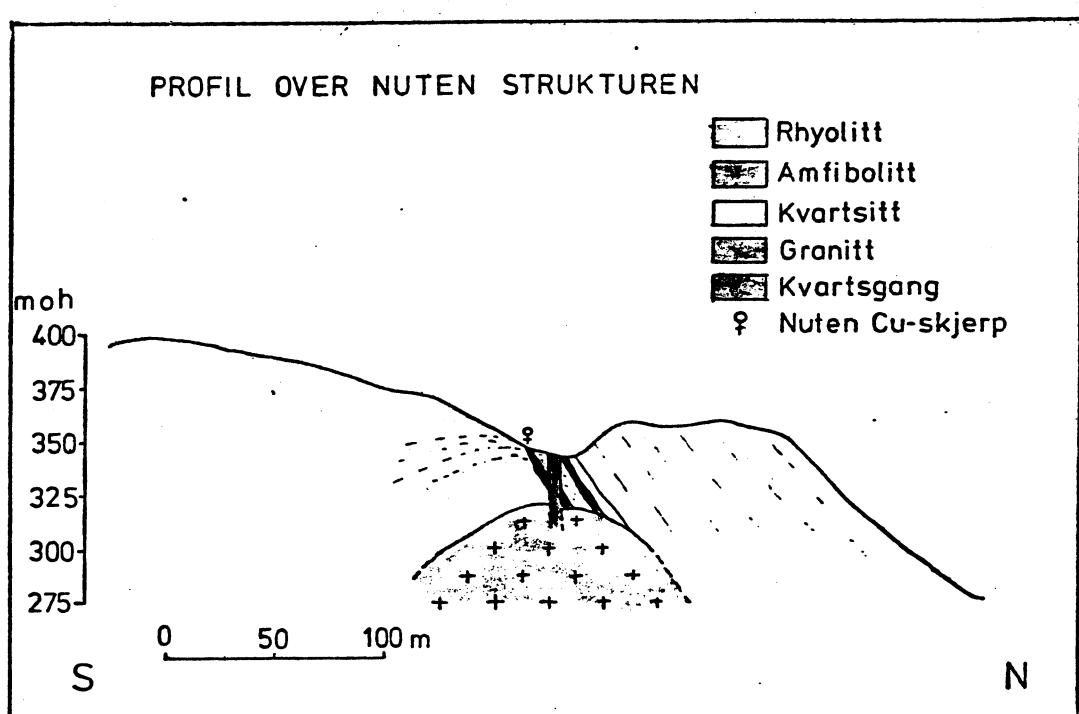
Nuten Cu-Bi skjerp.

Beliggenhed.

Nuten Cu-Bi skjerp ligger ca. 500 m SØ for Neshaug gård ved Hjartsjå på kartblad Flatdal (1614 III) i 1:50000, UTM-koordinater (8635,0725). Yderligere 500 m SØ for Nuten er registreret et skjerp (Bergstul), som har været skjerpet på Cu. Bergstul skjerp blev ikke lokaliseret, men ved lokaliteten UTM (8665, 0725)(se kort) blev observeret en mineraliseret kvartsåre med orienteringen N $50^{\circ}/85^{\circ}$, som blev forkastet af en mineraliseret sprække med orienteringen N $15^{\circ}/15^{\circ}$. Denne lokalitet formodes at ækvivalere Bergstul efter de beskrivelser der forelå. Nuten skjerp nås ved at dreje til venstre ad vejen ved savværket ca. 200 m øst for Hjartsjå gård i den østlige ende af Hjartsjå. Derefter fortsættes 200 m ad vejen til den deler sig. Ad den vej der går til højre fortsættes 1350 m til vejen i 350 moh deler sig : I T-krydsets SV-lige hjørne findes skjerpet, som måler 2x3 m i overfladen og 1-1½ m i dybden.

Geologisk beskrivelse.

Nuten Cu-Bi skjerp sidder i en 0,5-3,5 m tyk kvartsgang 350 moh i et topografisk markeret skar med samme orientering som kvartsgangen : N 82° (se kort). Området består geologisk af metasedimentære bjergarter tilhørende Rjukan og Seljord gruppen (skillelinien mellem de to grupper formodes at ligge indenfor de samme horisonter som kvartsgangen er afsat i). Stratigrafisk nedefra og op er der i området først afsat (se kort og skitse) en rhyolitisk bjergart derpå er afsat 0,1-1 m tyk basalt. Denne er overlejret af 1-2 m tyk rhyolit, hvorpå er afsat en ny



basaltisk lava, denne er i toppen brecciert af en ca. 0,5 m tyk rhyolit som i kontakten til de basaltiske xenolither får et storkornet amfibolitisk udseende, idet rhyoliten assimilerer basaltisk materiale. Denne

extrusive lagpakke formodes at tilhøre Rjukan gruppen og er overlejret af en tæt grå massiv kvartsit tilhørende Seljord gruppen. Lagpakken stryger N 50° - 60° og hælder 40° - 35° mod nord. Den rhyolitiske bjergart afviger lokalt fra denne orientering, på grund af dens fluktuerende flydebånding (muligheden af at bjergarten også er småfoldet kan ikke afdøves).

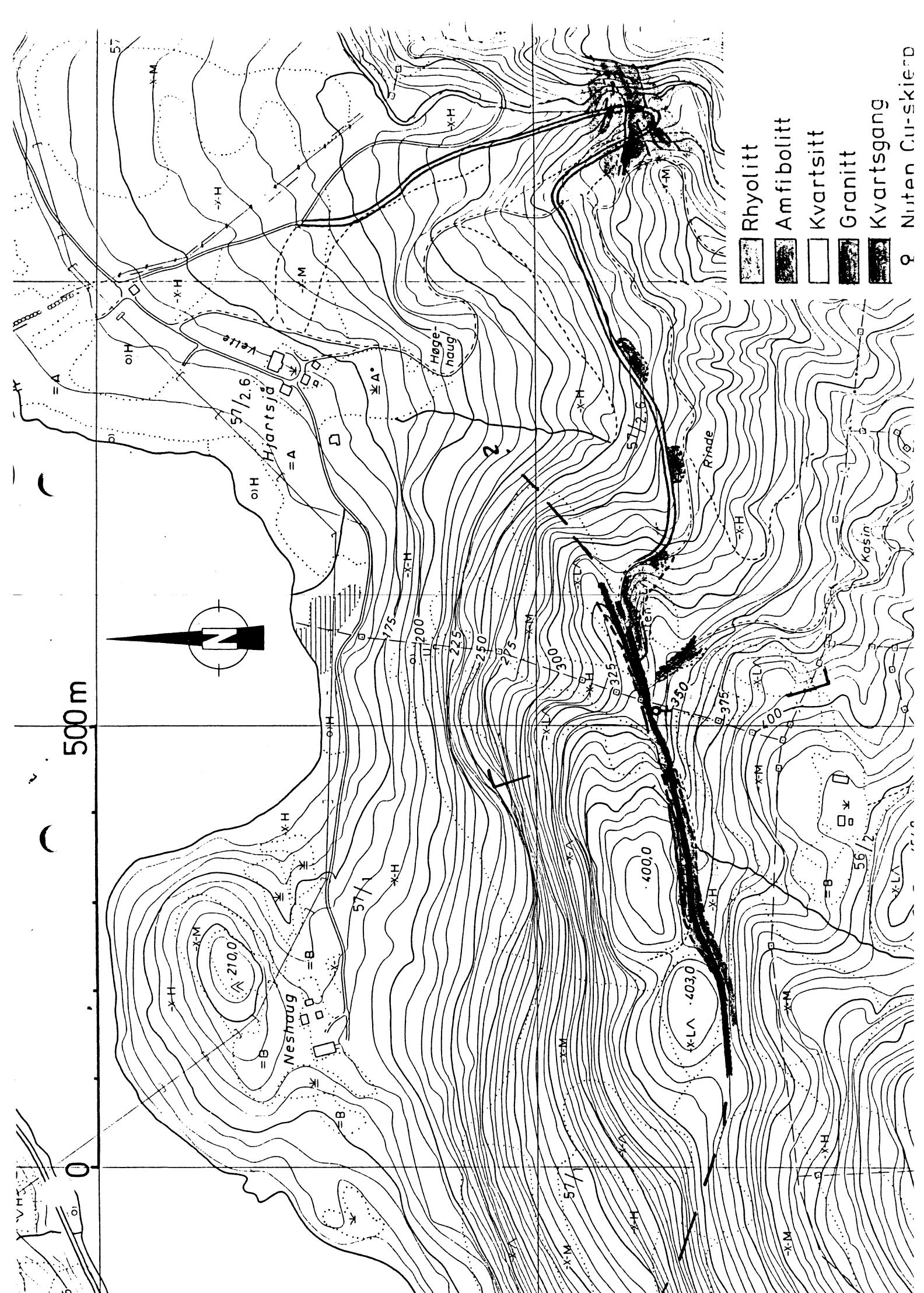
Denne metasedimentære lagpakke er først intruderet af en granitisk bjergart, som indenfor det undersøgte område er set med både vandrette og lodrette kontakter til de omgivende bjergarter. Denne mellemkornede rød/grønplette granit formodes derfor at udgøre toppen af et granitisk legeme. Op gennem disse bjergarter er intruderet en 0,5-3,5 m tyk kvartsgang som stryger N 82° og står subvertikalt. Kontakten til de omgivende bjergarter er ikke skarpt diskordant, men nærmest migmatitisk af udseende. De omgivende bjergarter er stærkt omdannede og shearede i kontakten til kvartsgangen. Sålede optræder graniten som grøn augengnejs (211o313) med chlorit langs shearflader i en 1-2 m tyk zone nær kvartsgangen, bjergarten er desuden silicificeret i kontakten (0,1-0,5 m til kvartsgangen). Den rhyolitiske bjergart er i lighed med graniten shearet og silicificeret i en zone nær kvartsgangen. Langs shear- og sprækkeflader er afsat chlorit og af og til ses også chalcopyrit.

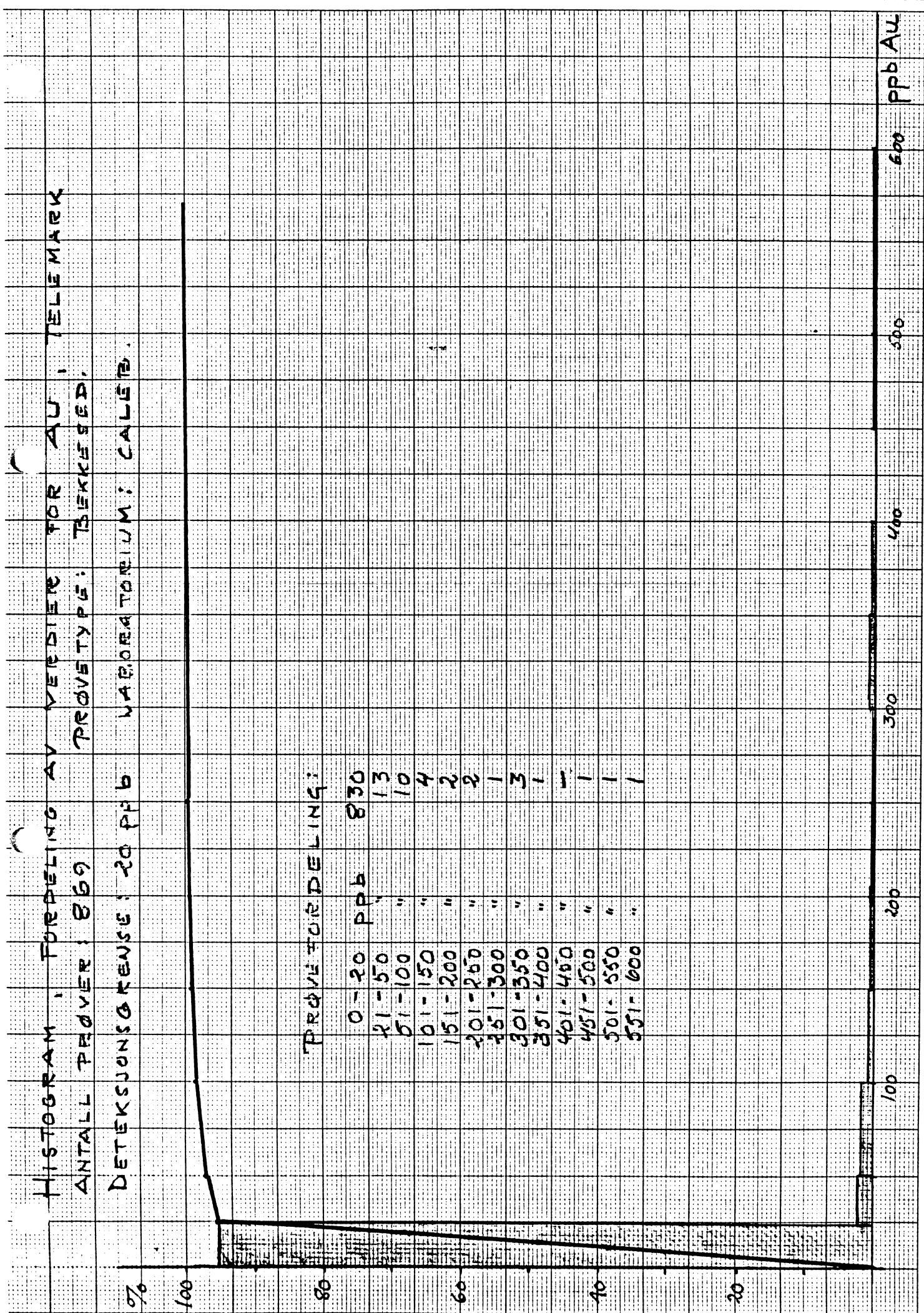
Mineralisering.

I forbindelse med kvartsgangen ses i området malmmineralerne pyrit, chalcopyrit og galena (bismuthholdig? Se 211o334). Chalcopyrit og pyrit er set sammen med chlorit som sprækkebelæg både i den shearede granit (se 211o310) og i den shearede rhyolit. Men langt den overvejende del af mineraliseringen findes i kvartsgangen. Årsagen til de rige pyrit- og chalcopyrit partier i kvartsgangen skal muligvis søges i den svage diskordans mellem kvartsåren og de amfibolitiske lag, hvorved der over en strækning på over 200 m sker en kraftig assimilering af basaltisk materiale. Dette medfører kraftig chloritisering samt pegmatitisk udvikling af det assimilerede materiale, ligesom dette fænomen også kan ses langs kontakten mellem kvartsgangen og graniten. Samtidig sker en koncentrering af især pyrit og chalcopyrit og muligvis også galena (bismuthholdig?). Kvartsgangen sensu stricto optræder derimod altid med mælkehvid kvarts og små miarolitiske hulrum med pyrit.

Udført arbejde.

Der er i området omkring Nuten udført geologisk kortlægning (se kort) samt indsamlet 9 lithoprøver deraf 5 chipsamples af både kvartsgangen og sidestensomdannelser. I området ved Bergstul er indsamlet 2 mineraliserede lithoprøver. Ved Nuten er desuden indsamlet 2 geokemiske prøver fra bække, som drænerer kvartsgangen i hele dens længde. De 2 prøver udgøres af et bækkesediment og en vaskeprøve.





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3	74750	64202130603	26.00	
4	74750	64202130604	16.00	
5	74750	64202130605	16.00	
6	74750	55202130606	16.00	
7	74750	55202130607	16.00	
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10	75750	57202130610	16.00	
11	75750	57502130611	16.00	
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13	74740	52402130613	16.00	
14	74710	55702130614	16.00	
15	74710	58202130615	16.00	
16	73740	51002130616	16.00	
17	73740	57202130617	16.00	
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19	78040	95602130619	16.00	
20	78170	95602130620	207.00	
21	78150	92802130621	10.00	
22	78150	1350622	10.00	
23	78150	109202130623	10.00	
24	78220	103402130624	10.00	
25	79220	103602130625	114.00	
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27	78220	101002130627	10.00	
28	78280	101502130628	10.00	
29	77920	110202130629	10.00	
30	78020	110902130630	24.00	
31	78120	105602130631	10.00	
32	78400	101602130632	10.00	
33	79240	105202130633	10.00	
34	79240	106202130624	10.00	
35	79420	111102130625	10.00	
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Year	Population (in millions)
1950	3.6
1955	3.9
1960	4.2
1965	4.5
1970	4.8
1975	5.1
1980	5.4
1985	5.7
1990	6.0
1995	6.3
2000	6.6
2005	6.9
2010	7.2
2015	7.5
2020	7.8
2025	8.1
2030	8.4
2035	8.7
2040	9.0
2045	9.3
2050	9.6
2055	9.9
2060	10.2
2065	10.5
2070	10.8
2075	11.1
2080	11.4
2085	11.7
2090	12.0
2095	12.3

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381	80700	17200	2120091
382	85700	17100	2120092
383	82100	16600	2120092
384	15000	16700	2120094
385	32740	18200	2130095
386	83150	17200	2120096
387	13325	12520	2120097
388	35340	15290	2120098
389	2421	12160	2120099
390	34270	17350	2130099
391	84260	18600	2130091
392	84380	15930	2120092
393	24500	20450	2120093
394	34400	1540	2130094
395	23760	20080	2130095
396	62260	15320	2120096
397	52210	19230	2120097
398	32270	22120	2130098
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401	8670	2070	2120101
402	81600	1940	2120102
403	81150	1940	2130103
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406	85780	32230	2120106
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409	86930	27200	2131009
410	60240	27340	2120101
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412	60250	27300	2131012
413	36900	2610	2131013
414	59320	25910	2121014
415	86950	25910	2121015
416	26950	26480	2121016
417	58450	34780	2131017
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422	62700	32780	2131022
423	82700	22670	2121023
424	70380	19960	2131024
425	75900	32960	2121025
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427	63600	31100	2131027
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429	86680	30600	2131029
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431	87600	32150	2121031
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433	82100	32500	2121033
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675	82080	17500	2121275	10.00
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678	87480	2500	2121278	10.00
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717	0	79450	19550	2131317	10.00
718	0	79760	19440	2131318	10.00
719	0	79920	19200	2131319	10.00
720	0	79560	19160	2131320	10.00
721	0	79520	18920	2131321	10.00
722	0	72400	19220	2131322	10.00
723	0	77550	26540	2131323	10.00
724	0	77190	22450	2131324	10.00
725	0	77130	23400	2131325	10.00
726	0	77550	23760	2131326	10.00
727	0	77500	21280	2131327	10.00
728	0	72400	20300	2131328	10.00
729	0	72330	20500	2131329	10.00
730	0	72130	2131330	10.00	
731	0	20900	20600	2131331	10.00
732	0	61790	28000	2131332	10.00
733	0	72720	30910	2131333	10.00
734	0	73600	24690	2131338	10.00
735	0	72260	24150	2131339	10.00
740	0	72260	22450	2131340	10.00
741	0	72170	23530	2131341	10.00
742	0	75730	30000	2131342	10.00
743	0	76020	20060	2131343	10.00
744	0	75620	30240	2131344	10.00
745	0	75550	24190	2131345	10.00
746	0	73360	23860	2131346	10.00
747	0	73350	23670	2131337	10.00
748	0	72250	24690	2131338	10.00
749	0	72260	22450	2131339	10.00
750	0	72170	23530	2131340	10.00
751	0	75340	30330	2131341	10.00
752	0	76660	27300	2131346	10.00
747	0	74920	22350	2131347	10.00
748	0	72940	23200	2131348	10.00
749	0	73760	23450	2131349	10.00
750	0	59200	25500	2131350	10.00
751	0	77170	6130	2131351	10.00
752	0	76550	6330	2131352	10.00
753	0	76550	6430	2131353	10.00
754	0	76550	6510	2131354	10.00
755	0	78330	14150	2131355	10.00
756	0	77540	14520	2131356	10.00

757	60	10.00	2131357
758	7.410	14.00	2131358
759	7.420	15.00	2131359
760	7.5420	14.520	2131360
761	7.5600	15.340	2131361
762	7.6280	12.640	2131362
763	8.9460	24.750	2131363
764	7.690	24.530	2131364
765	7.710	24.700	2131365
766	7.7170	25.570	2131366
767	7.6710	28.160	2131367
768	8.9280	28.660	2131368
769	8.420	32.700	2131369
770	8.4510	33.200	2131370
771	8.2660	34.340	2131371
772	8.2690	34.350	2131372
773	7.2450	35.900	2131373
774	7.3550	34.500	2131374
775	7.150	35.400	2131375
776	7.0720	35.475	2131376
777	7.9650	22.170	2121377
778	8.9350	21.300	2131378
779	8.1330	21.980	2131379
780	7.5270	21.560	2121380
781	7.8690	21.660	2131381
782	7.8540	21.560	2131382
783	7.620	21.960	2131383
784	7.7070	21.500	2131384
785	7.6380	21.630	2131385
786	7.5760	20.660	2131386
787	7.5550	20.620	2121387
788	7.5750	21.120	2131388
789	7.5980	21.440	2131389
790	7.5780	21.700	2131390
791	7.5240	21.350	2131391
792	7.4500	21.760	2121392
793	7.5750	22.110	2121393
794	7.7790	21.200	2131394
795	7.5750	21.500	2131395
796	7.4730	21.430	2131396
797	7.4100	21.620	2131397
798	7.4220	21.320	2131398
799	7.3650	20.620	2121399
800	7.2900	20.780	2131400
801	7.2590	22.440	2131401
802	7.3180	21.320	2131402
803	7.170	22.40	2131403
804	75.430	22.750	2131404
805	74.330	22.750	2131405
806	81.530	27.600	2131406
807	81.480	27.120	2131407
808	81.70	27.670	2131408
809	82.110	25.100	2131409
810	81.675	30.510	2131410
811	81.200	30.000	2131411
812	82.990	29.410	2131412
813	81.610	30.780	2131413
814	81.000	35.620	2131414
815	82.770	31.660	2131415
816	81.9	32.620	2131416
817	83.150	35.140	2131417
818	78.750	34.200	2131418
819	80.500	31.750	2131419

320	0	31000	2131420	10.00	
321	0	34750	2131421	10.00	
322	0	21730	2131422	10.00	
323	0	32230	2131423	10.00	
324	0	35180	7200	10.00	
325	0	35200	7540	10.00	
326	0	35260	7630	226.00	
327	0	29260	7580	2131427	10.00
328	0	87230	7420	2131428	10.00
329	0	37100	7180	2131429	10.00
330	0	37100	7180	2131430	10.00
331	0	36120	8250	2131431	10.00
332	0	36340	9220	2131432	10.00
333	0	72650	37200	2131433	10.00
334	0	74500	36600	2131434	10.00
335	0	72600	3200	2131435	10.00
336	0	73900	24450	2131436	10.00
337	0	74100	29450	2131437	10.00
338	0	74350	27500	2131438	10.00
339	0	76050	37050	2131439	10.00
340	0	76600	36850	2131440	10.00
341	0	74750	37150	2131441	10.00
342	0	75300	37150	2131442	10.00
343	0	74700	27050	2131443	10.00
344	0	76000	36850	2131444	10.00
345	0	76500	36900	2131445	10.00
346	0	77450	35900	2131446	10.00
347	0	77600	35900	2131447	10.00
348	0	77700	27700	2131448	10.00
349	0	77850	37050	2131449	10.00
350	0	77950	27250	2131450	10.00
351	0	77650	35600	2131451	10.00
352	0	79250	34900	2131452	10.00
353	0	79250	34900	2131453	10.00
354	0	81900	37600	2131454	10.00
355	0	81900	37600	2131455	10.00
356	0	82300	37500	2131456	10.00
357	0	82400	37850	2131457	10.00
358	0	84100	37250	2131458	10.00
359	0	84700	37350	2131459	10.00
360	0	85200	37200	2131460	10.00
361	0	87700	32250	2131461	10.00
362	0	86700	32250	2131462	10.00
363	0	87100	32500	2131463	10.00
364	0	85900	37500	2131464	10.00
365	0	885	37200	2131465	10.00
366	0	87400	37400	2131466	10.00

EXHIBIT MAP - CONTENTS:

SAMPLE MAP 1 - Svartdal BP037-5-1.
Samples represented - 2110043-46, 60.

SAMPLE MAP 2 - Ambjørndalen BP037-5-2.
Samples represented - 2110001-2, 4-9, 13-14,
47-59, 61-75, 77-82, 89, 146-148, 226.

SAMPLE MAP 3 - Nutheim BP037-5-3.
Samples represented - 2110076.

SAMPLE MAP 4 - Ståldalen BP038-5-1.
Samples represented - 2110223.

SAMPLE MAP 5 - Barstadhovi BP038-5-3.
Samples represented - 2110010-12, 18-42, 83-87,
91-145, 201-222, 227-255.

SAMPLE MAP 6 - Hovdejord BP038-5-4.
Samples represented - 2110088, 225.

SAMPLE MAP 7 - Svinom BP039-5-3.
Samples represented - 2110224.

SAMPLE MAP 8 - Bisminuten BQ038-5-1.
Samples represented - 2110017.

SAMPLE MAP 9 - Hjartdal BQ038-5-3.
Samples represented - 2110016.

SAMPLE MAP 10 - Gvammen BQ038-5-4.
Samples represented - 2110003.

SAMPLE MAP 11 - Kivikdalen BQ039-5-3.
Samples represented - 2110015.

SAMPLE MAP 12 - "Flatdal"
Samples represented - 2110090.